

Supporting Information for “Contrasting regional carbon cycle responses to seasonal climate anomalies across the east-west divide of temperate North America”

B. Byrne¹, J. Liu^{1,2}, A. A. Bloom¹, K. W. Bowman^{1,3}, Z. Butterfield⁴,

J. Joiner⁵, T. F. Keenan^{6,7}, G. Keppel-Aleks⁴, N. C. Parazoo¹, and Y. Yin²

¹Jet Propulsion Laboratory, California Institute of Technology, CA, USA

²Division of Geological and Planetary Sciences, California Institute of Technology, Pasadena, CA, USA

³Joint Institute for Regional Earth System Science and Engineering, University of California, Los Angeles, USA

⁴Department of Climate and Space Sciences and Engineering, University of Michigan, Ann Arbor, MI, USA

⁵Laboratory of Atmospheric Chemistry and Dynamics, NASA Goddard Space Flight Center, Greenbelt, MD 20771, USA

⁶Earth and Environmental Sciences Area, Lawrence Berkeley National Laboratory, Berkeley, California, USA

⁷Department of Environmental Science, Policy and Management, University of California, Berkeley, Berkeley, California, USA

Contents of this file

1. Text S1
2. Figures S1 to S9
3. Table S1

Text S1. Flux inversion NEE

CT2017 (Peters et al. (2007), with updates documented at <https://www.esrl.noaa.gov/gmd/ccgg/carbontracker/>) optimizes NEE by assimilating flask and in situ CO₂ measurements. It employs an ensemble Kalman filter approach to assimilate CO₂ with atmospheric chemical transport simulated by the TM5 offline atmospheric model (Krol et al., 2005). For CT2017, TM5 is driven by ERA-Interim assimilated meteorology from the European Centre for Medium-Range Weather Forecasts (ECMWF), with a horizontal resolution of 3° × 2° globally and 1° × 1° in a nested grid over temperate North America.

CT-L provides estimates of posterior NEE over temperate North America by performing analytic Bayesian inversions assimilating flask and in situ CO₂ measurements (Hu et al., 2019). For these inversions, contributions to the observations from background CO₂ inflow over temperate North America and surface CO₂ emissions from fossil fuel and fire emissions are pre-subtracted from the observations. Footprint sensitivities of the observations to surface fluxes are calculated using the high-resolution Weather Research and Forecasting–Stochastic Time–Inverted Lagrangian Transport (WRF-STILT) model at 10 km spatial resolution over the domain of interest, providing footprints at 1x1 and hourly temporal resolution and represent simulated upwind influences over 10 days before each measurement. An ensemble of 18 flux inversions is performed by varying prior NEE and boundary conditions in the inversion. Here we examine the ensemble mean from CT-L. Monthly NEE fluxes at 1x1 spatial resolution were downloaded from <https://doi.org/10.15138/3dw1-5c37>.

The CAMS greenhouse gases inversion system (v18r3) (Chevallier et al., 2005, 2010; Chevallier, 2013) assimilates surface air-sample CO₂ measurements from 129 sites over the globe. 4-DVar is employed to optimize day-time and night-time NEE at 8-day temporal resolution of 1.875°×3.75° model grid. Tracer transport is performed using the Laboratoire de Météorologie Dynamique (LMDz) general circulation model version LMDz6A (Remaud et al., 2018). These data were downloaded from <https://atmosphere.copernicus.eu/>.

References

- Byrne, B., Liu, J., Lee, M., Baker, I. T., Bowman, K. W., Deutscher, N. M., ... Wunch, D. (2020). Improved constraints on northern extratropical CO₂ fluxes obtained by combining surface-based and space-based atmospheric CO₂ measurements. *Journal of Geophysical Research: Atmospheres*, 125. doi: 10.1029/2019JD032029
- Chevallier, F. (2013). On the parallelization of atmospheric inversions of CO₂ surface fluxes within a variational framework. *Geoscientific Model Development*, 6(3), 783–790. Retrieved from <https://gmd.copernicus.org/articles/6/783/2013/> doi: 10.5194/gmd-6-783-2013
- Chevallier, F., Ciais, P., Conway, T., Aalto, T., Anderson, B., Bousquet, P., ... Worthy, D. (2010). CO₂ surface fluxes at grid point scale estimated from a global 21 year reanalysis of atmospheric measurements. *Journal of Geophysical Research: Atmospheres*, 115(D21).
- Chevallier, F., Fisher, M., Peylin, P., Serrar, S., Bousquet, P., Bréon, F.-M., ... Ciais, P. (2005). Inferring CO₂ sources and sinks from satellite observations: Method and application to TOVS data. *Journal of Geophysical Research: Atmospheres*, 110(D24).

- Hu, L., Andrews, A. E., Thoning, K. W., Sweeney, C., Miller, J. B., Michalak, A. M., ... others (2019). Enhanced north american carbon uptake associated with el niño. *Science advances*, 5(6), eaaw0076.
- Krol, M., Houweling, S., Bregman, B., Broek, M., Segers, A., Velthoven, P. v., ... Bergamaschi, P. (2005). The two-way nested global chemistry-transport zoom model tm5: algorithm and applications. *Atmos. Chem. Phys.*, 5(2), 417–432. doi: 10.5194/acp-5-417-2005
- Peters, W., Jacobson, A. R., Sweeney, C., Andrews, A. E., Conway, T. J., Masarie, K., ... others (2007). An atmospheric perspective on North American carbon dioxide exchange: CarbonTracker. *Proc. Natl. Acad. Sci.*, 104(48), 18925–18930. doi: 10.1073/pnas.0708986104
- Remaud, M., Chevallier, F., Cozic, A., Lin, X., & Bousquet, P. (2018). On the impact of recent developments of the lmdz atmospheric general circulation model on the simulation of CO₂ transport. *Geoscientific Model Development*, 11(11), 4489–4513. Retrieved from <https://gmd.copernicus.org/articles/11/4489/2018/> doi: 10.5194/gmd-11-4489-2018

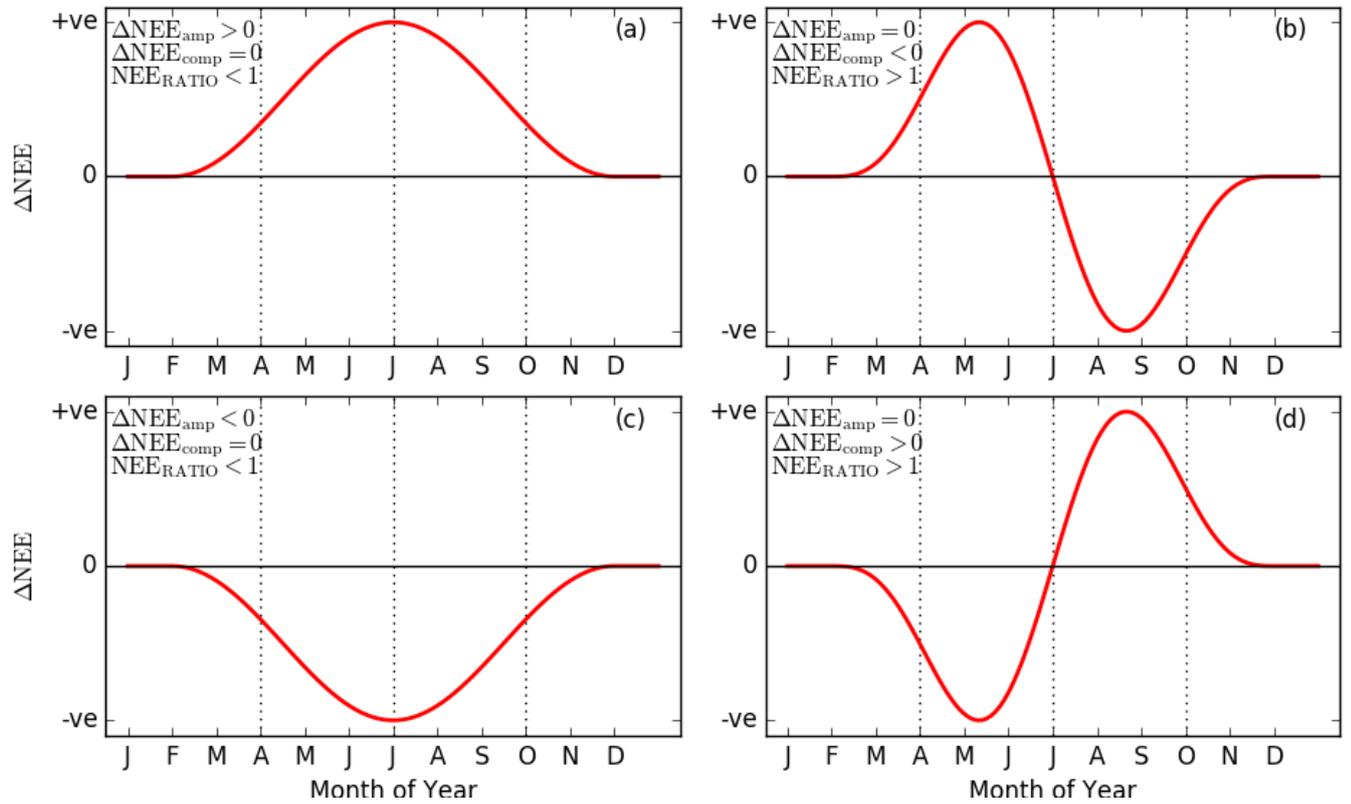


Figure S1. Illustration of amplification and compensation for NEE. (a) Positive amplification with no compensation, (b) no amplification with negative compensation, (c) negative amplification with no compensation, and (d) no amplification with positive amplification. Note that the same scheme is applied to GPP.

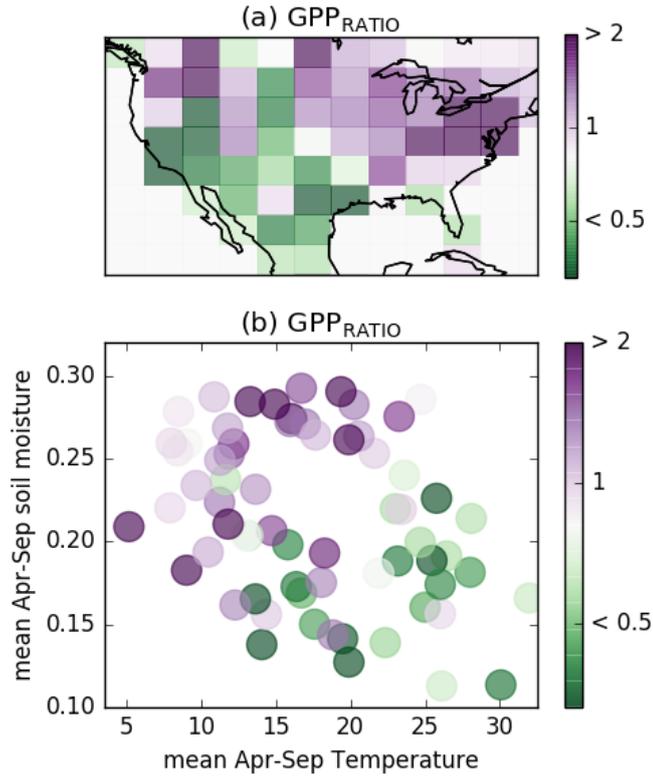


Figure S2. Relative magnitudes of seasonal compensation and amplification. (a) GPP_{RATIO} over 2010–2015 at $4^\circ \times 5^\circ$. (b) GPP_{RATIO} plotted as a function of Apr-Sep mean soil temperature (K) and soil moisture ($m^3 m^{-3}$).

Table S1. Correlation coefficient (R) between CO_2 flux anomalies and $\Delta T_{Apr-Sep}$, $\Delta M_{Apr-Sep}$, $\Delta P_{Apr-Sep}$, or $\Delta TWS_{Apr-Sep}$. Correlations cover the period 2001–2017 for ΔGPP and 2010–2015 for ΔNEE , except for correlations with ΔTWS which cover 2003–2014 and 2010–2014. For eastern North America, 2012 is excluded for ΔGPP correlations because it is an extreme event and has a large impact on the correlation.

Region	Environ Var	ΔGPP_{amp}	ΔGPP_{comp}	ΔNEE_{amp}	ΔNEE_{comp}
West	$\Delta T_{Apr-Sep}$	-0.71	-0.41	0.63	0.22
West	$\Delta M_{Apr-Sep}$	0.91	0.09	-0.66	-0.13
West	$\Delta P_{Apr-Sep}$	0.78	0.31	-0.47	-0.21
West	$\Delta TWS_{Apr-Sep}$	0.50	0.20	-0.70	-0.19
East	$\Delta T_{Apr-Sep}$	-0.09	-0.81	-0.46	0.89
East	$\Delta M_{Apr-Sep}$	0.72	0.41	0.78	-0.49
East	$\Delta P_{Apr-Sep}$	0.35	0.31	0.48	-0.44
East	$\Delta TWS_{Apr-Sep}$	0.56	0.31	0.81	-0.30

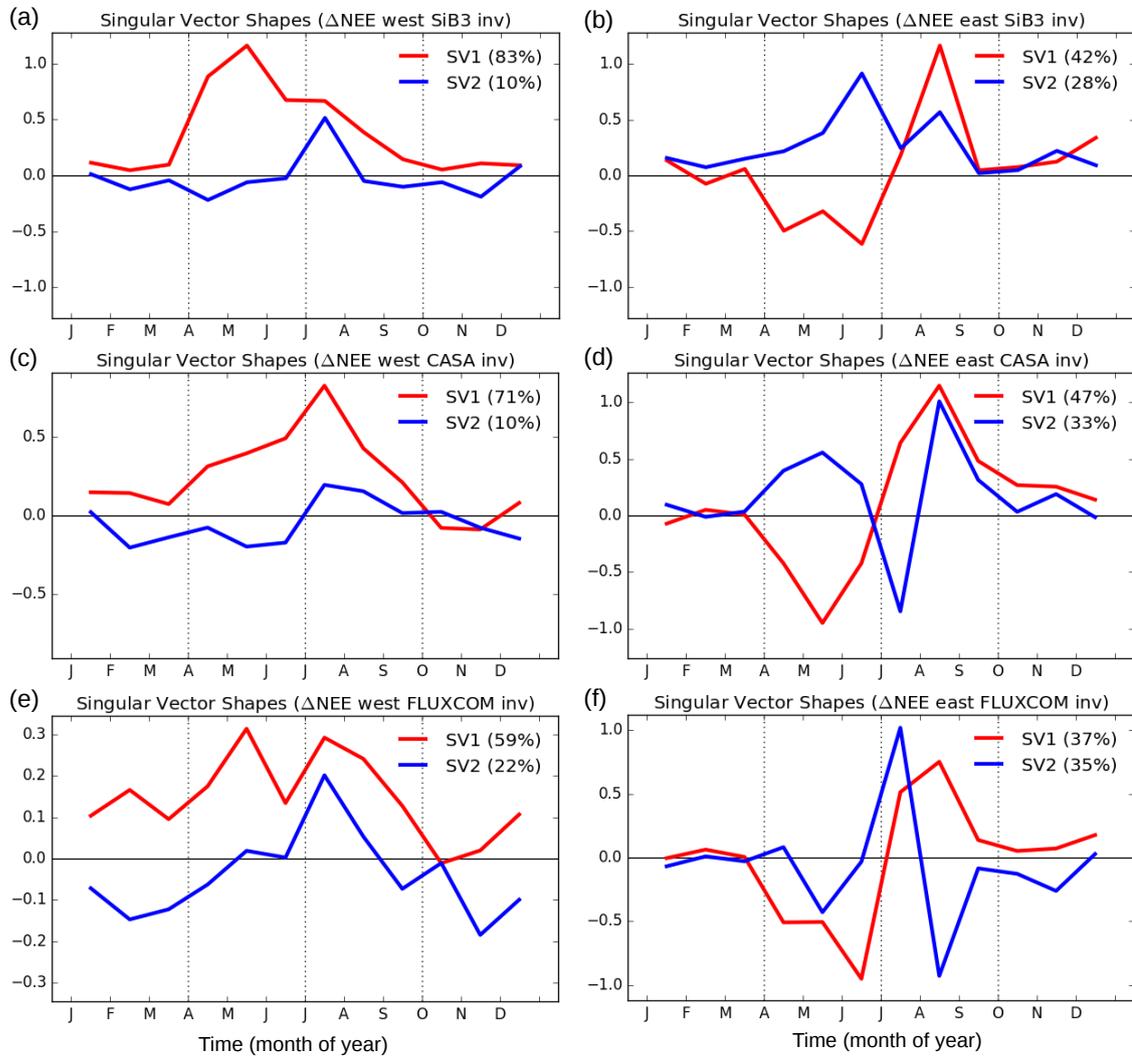


Figure S3. SVD analysis for individual flux inversions from Byrne et al. (2020) for (a) western and (b) eastern North America

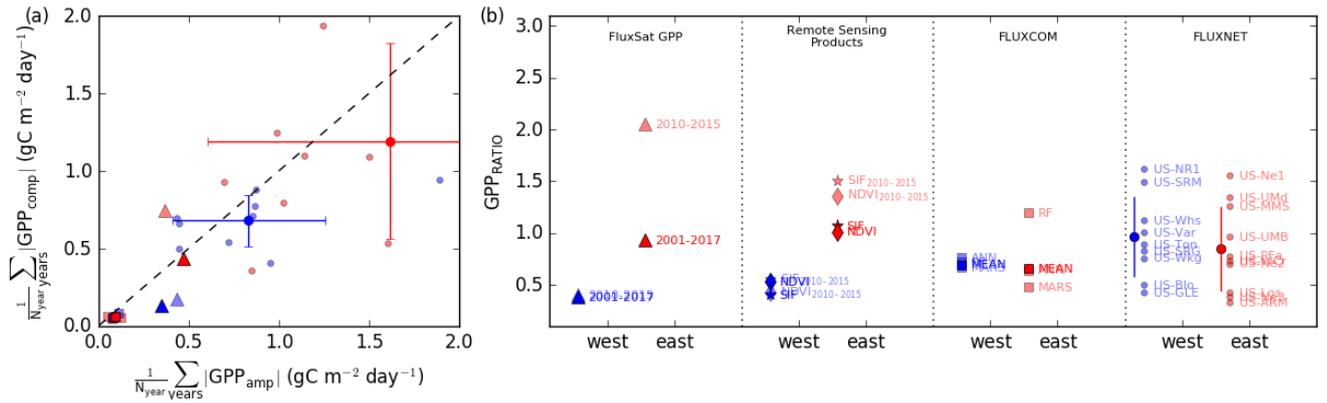


Figure S4. (a) Mean magnitude of GPP compensation versus mean magnitude of GPP amplification across multiple years. (b) GPP_{RATIO} over eastern and western North America for (left-to-right) FluxSat, remotely sensed products (GOME-2 SIF and MODIS NDVI), FLUXCOM GPP, and FLUXNET sites with 6+ years of data within the eastern and western domains. Partially transparent symbols show values over 2010–2015 and solid colors are for the entire time period examined in this study for a given dataset. Note that NDVI and SIF are not shown in panel (a) because they do not have units of GPP.

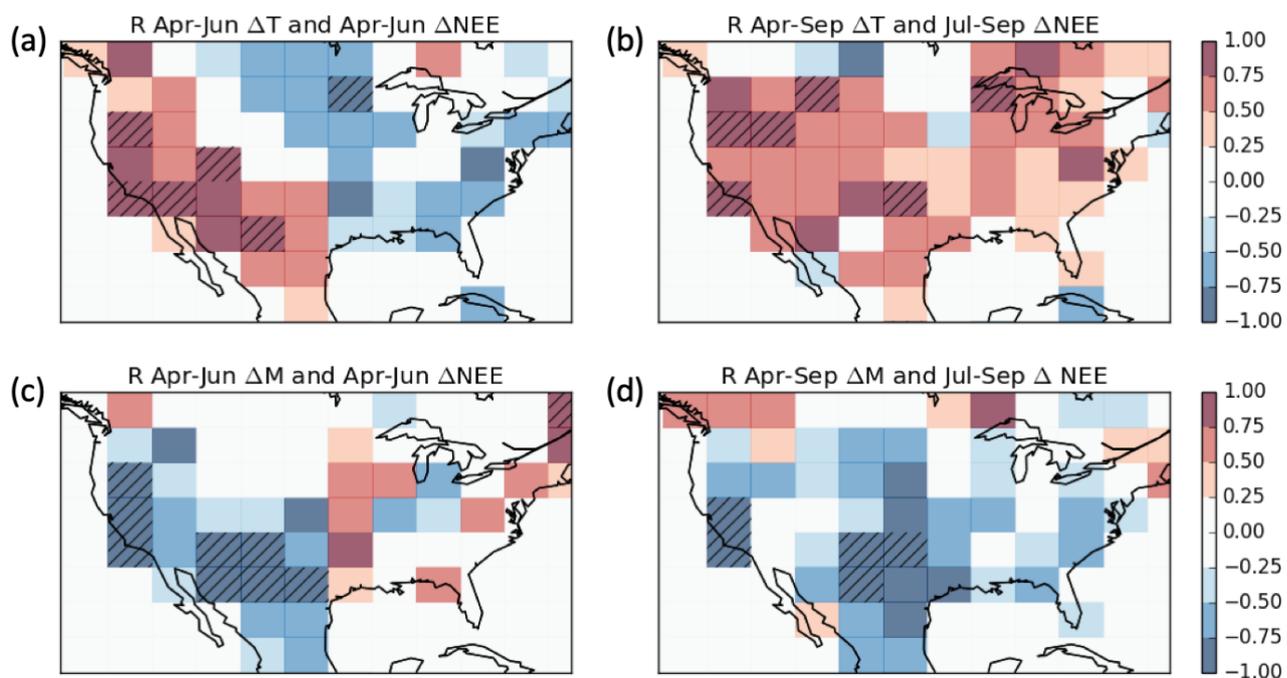


Figure S5. Relationship between ΔNEE and variations in climate. Coefficient of correlation (R) over 2010–2015 for $4^\circ \times 5^\circ$ grid cells between (a) Apr–Jun ΔT and Apr–Jun ΔNEE , (b) Apr–Sep ΔT and Jul–Sep ΔNEE , (c) Apr–Jun ΔM and Apr–Jun ΔNEE and (d) Apr–Sep ΔM and Jul–Sep ΔNEE . Hatching shows grid cells for which $P < 0.05$.

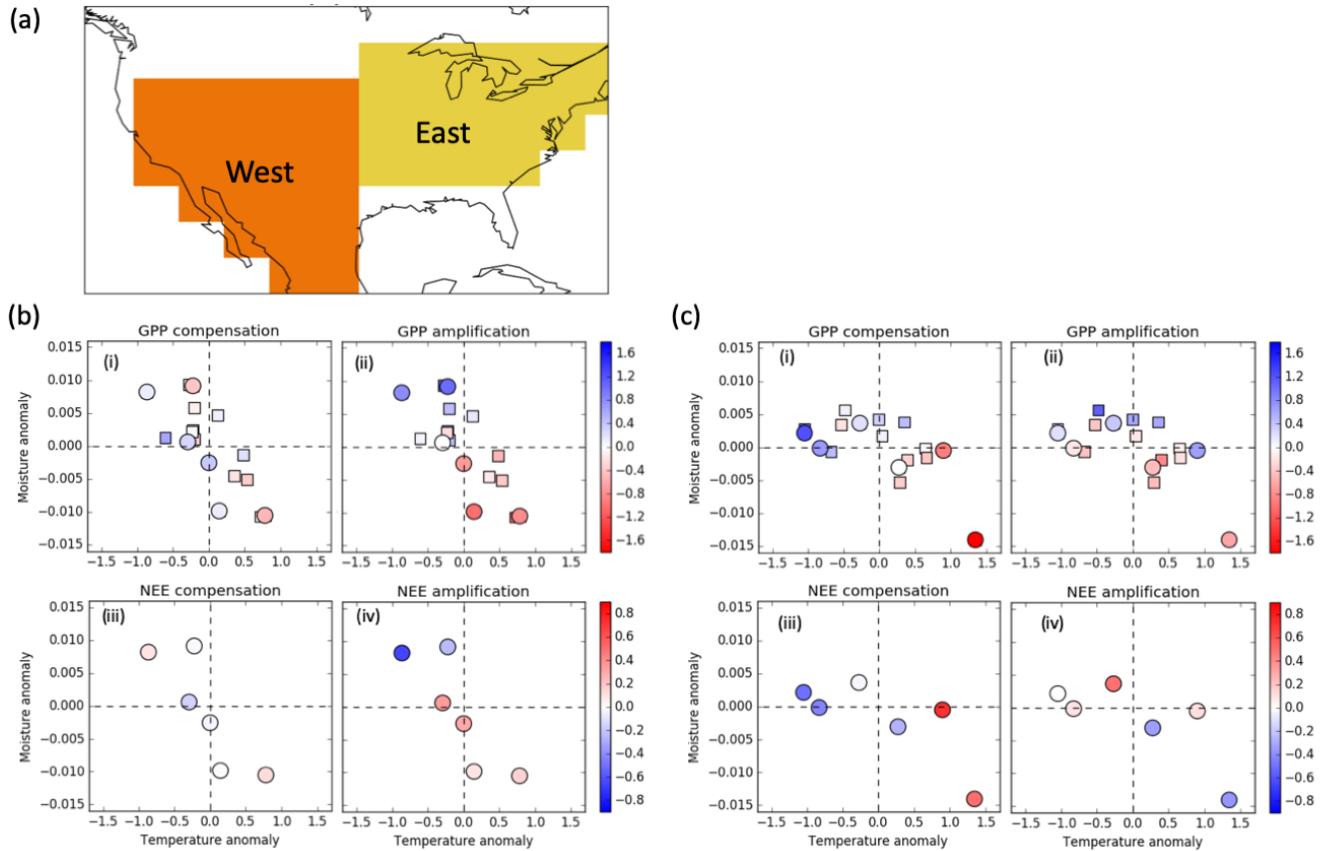


Figure S6. Regional seasonal compensation and amplification components as a function of ΔT and ΔM . (a) Spatial extent of western (orange) and eastern (yellow) regions of North America. (b) Western and (c) Eastern scatter plots of (i) $\Delta \text{GPP}_{\text{comp}}$, (ii) $\Delta \text{GPP}_{\text{amp}}$, (iii) $\Delta \text{NEE}_{\text{comp}}$ and (iv) $\Delta \text{NEE}_{\text{amp}}$ as a functions of ΔT and ΔM . NEE covers the period 2010–2015 and GPP covers the period 2001–2017, with circles indicating points over 2010–2015 and squares indicating points outside this time period. Colorbars have units of PgC yr^{-1} .

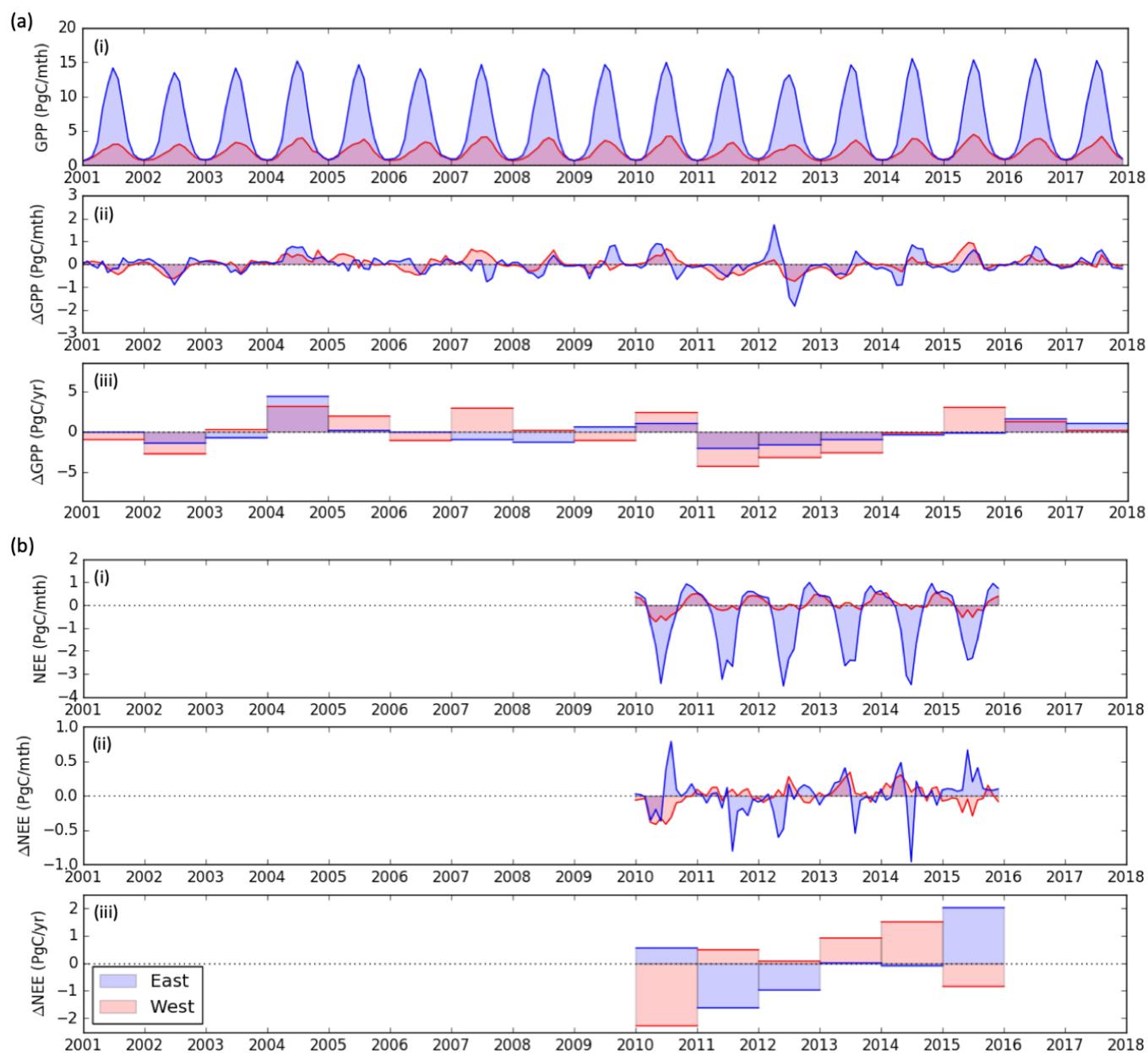


Figure S7. Timeseries of (a) GPP (2001–2017) and (b) NEE (2010–2015) in western (shaded red area) and eastern (shaded blue area) North America. For GPP and NEE, panel (i) shows the seasonal cycle, (ii) shows the monthly anomalies, and (iii) shows the yearly anomalies.

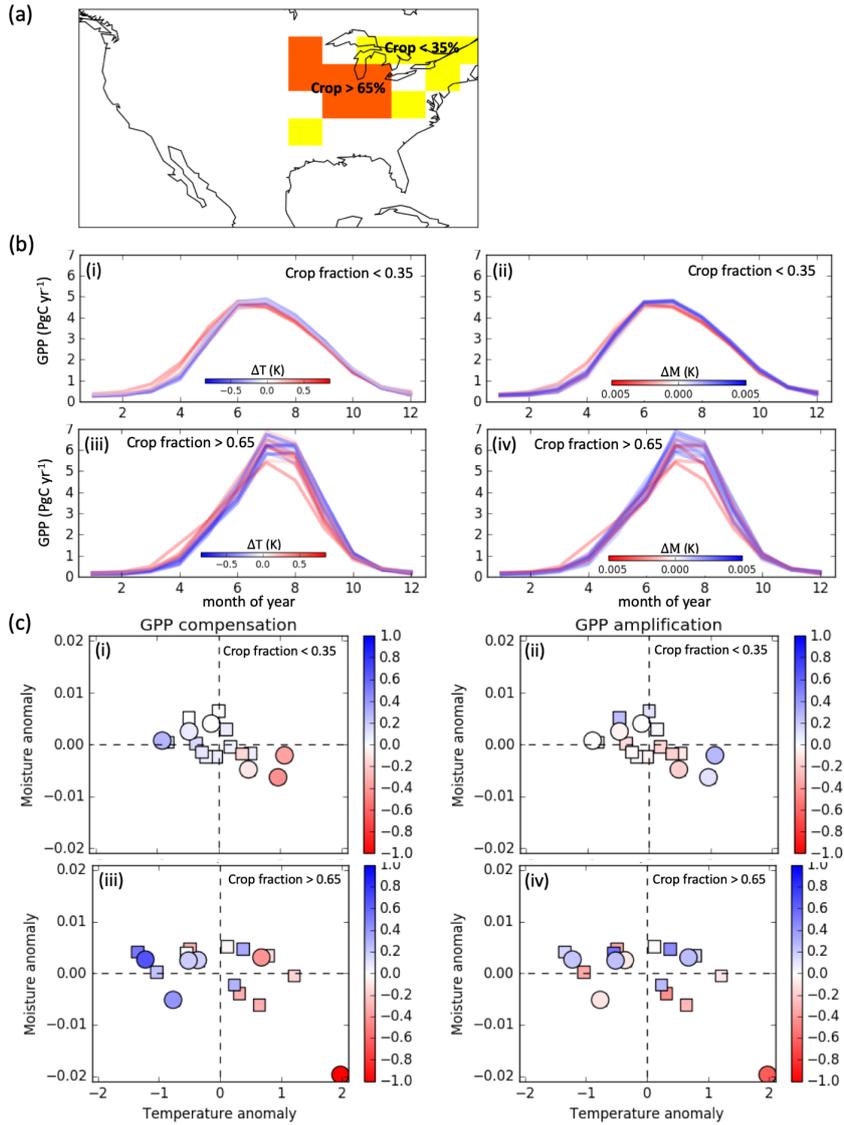


Figure S8. (a) Grid cells with crop fractions $> 65\%$ or $< 35\%$. (b) Time series of 2001-2017 GPP as a function of month of year for grid cells with crop fractions $< 35\%$ ((i) and (ii)) and crop fractions $> 65\%$ ((iii) and (iv)). Curves are colored by Apr-Sep temperature anomaly for (i) and (iii), and are colored by Apr-Sep moisture anomaly for (ii) and (iv). (c) Seasonal compensation ((i) and (iii)) and amplification ((ii) and (iv)) components for grid cells with crop fractions $< 35\%$ and $> 65\%$. Note that the outlier for crop fractions $> 65\%$ is due to the 2012 North American Drought. Colorbars have units of PgC yr^{-1} .