

**Robust Earthquake Early Warning at a Fraction of the Cost: ASTUTI Costa Rica**

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Author Response to Peer Review Comments

**Peer Review Comments on 2021AV000407R**

**Reviewer #1**

The author provide a smartphone-based EEW system and show its performance in Costa Rica. What I am most interested in is the data that generates by smartphones (or precisely the MEMS sensor in the phone). As the author didn't mention the phone model that was used in this EEW system, readers may want to know more about the data situation. So I suggest you add some more details about the phone model selection, such as what's the considerations about the phone model? Did you take any tests on

different models (MEMS)? and what are the main differences? And I partly-disagree with the authors' view about "Smartphones used in a fixed network can provide earthquake early warning performance similar to scientific-grade instrumentation". The author indeed shows the reliability of this smartphone-based EEW system, but it will be more convincing if the author could provide the lower limit magnitude of this system.

*[Please see attachment that begins on the next page.]*

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## Reviewer #2

Review for AGU Advances

Smartphone-based Earthquake Early Warning with a Fixed Network: ASTUTI Costa Rica  
Brooks, B. et al.

I consider that the paper submitted by Brooks and his co-authors constitutes a novel and interesting idea that will be of interest to a broad range of readers in the Earth and social sciences communities. As the authors stress, and I fully agree with them, this technology, once perfected and more developed, may represent an interesting alternative to the deployment of Earthquake Early Warning Systems (EEMS) in countries that have neither the funding nor the communications infrastructure to install and operate an EEMS with more sophisticated instrumentation.

I believe the paper is very interesting and innovative and I would recommend its publication in AGU Advances. I would encourage the authors to consider some of the comments written below, as well as those that I added as corrections to the *Word* document attached to this review. I used this *Word* document because it was easier than to comment on the tightly spaced *pdf* pre-print available. Also, it has comments included by other reviewers (?). I hope that this marked document makes it easier for the authors to produce a revised manuscript.

Other comments are listed below:

1. The paper is extremely long. At times, it seems like the authors want to publish a review of EEMS. It is always important to provide the background, but the authors go to unnecessary lengths. I would encourage them to shorten some sections substantially, I elaborate below.
2. The paper seems to be written in haste and not thoroughly checked before submitting. Some sentences are awkward, confusing or simply unnecessary; I point out many of them in the marked document.
3. An example of this impression that it was rapidly put together is that in many cases the authors cite some references in the text with the full name and last name of the authors and not simply by last name as is the norm. This certainly gives the impression that the paper was not proofed critically by the authors before submitting.
4. I agree with the authors that for a country the size of Costa Rica, an EEMS that warns for every earthquake detected may be acceptable. However, what the authors call the "boy cry wolf" syndrome may wear down the population in time. Having worked in collaboration with the Mexican EWS, I know that the public is very demanding and unforgiving with systems like this. In addition, the authors only touch lightly on the presence of intermediate-depth and crustal earthquakes in Costa Rica. They admit that

ASTUTI may not work for them, but this is done only in passing. I would ask the authors to consider a future magnitude threshold and mention it at least in this initial paper.

5. There are some lengthy discussions that could be summarized in Tables for the benefit of the authors. I found it difficult and boring to read, all the statistics regarding the delay times for the earthquakes tested. Also, the discussion towards the end of the paper of percentages of the population that felt shaking or not and its relation to the DYFI data is extremely dry. I would encourage the authors to consider putting the data on Tables and refer the readers to them.

6. Following on the last comment, the Discussion section is too long and repetitive of what is said previously in other parts of the paper. Also, it makes the Conclusions section seem unnecessary and gratuitous. The authors know well that many scientists today read carefully the abstract and conclusions of papers. The authors should give a critical review of how this section is written today and shorten it, emphasizing their findings and innovations.

7. Like all EEWS, to become a useful tool for society, the public has to be educated in its use. I know this is no easy task. However, I believe the authors' definition of Missed Alerts covers too broad a range of earthquakes. I think the authors have to be realistic as to where earthquakes need to occur for ASTUTI to be an effective and useful tool. Whenever they roll it out, this has to be made clear. I encourage them to think about it and include it in the paper.

8. In terms of the False Alerts, the authors point out to power failures of the grid line; a common problem in many countries. Although the idea is to keep it as a low-cost system, an inexpensive power supply would solve this problem. It is not clear from their discussion how many times this happened.; it would be useful to know. I cannot help but to suggest to the authors also reference the statistics of SASMEX in terms of FA: one false alert during its operational life.

9. Some figures are confusing and of low quality. For example, Figure 7a shows both the average ASTUTI latency for all events (or at least this is what I understand of it) together with a discussion on the performance of the system during the 27 July 2020 earthquake. Figure 7a is not referred to in the text and the reader, as I did, is forced to speculate what it means and what it represents. If Figures 7b, 7c and 7d refer only to the 27 July quake and 7a reflects the median latency for all quakes they should not be in the same Figure. This needs to be corrected.

10. In Figure 8 what authors call copper population is confusing. There are too many color bars. The ones for DYFI are clear. The others, however, are confusing. Three color bars have the title Pop. What do they mean? I would encourage the authors to improve the figure making it more didactic and clarifying the caption.

11. The main question to me is whether Amazon Web Service Notification will work when the system is implemented fully, and thousands of users need to be notified (maybe a few million). The authors touch on this in passing. I admit to being totally ignorant of this web service, however, having witnessed how pilot programs using cell phone notification worked well for a reduced number of users but failed dramatically when the number of users increased, I have a grave concern whether this solution will work. Needless to say, if you do not notify broadly and on timely, the system is futile, no matter how sophisticated it may be technically. Not understanding how this web notification works, I would encourage the authors to include tests that confirm that it is a usable alternative.

12. Finally, the question of why the authors, particularly the local ones, do not mention the option to use the excellent national strong motion network data is puzzling. Costa Rica has perhaps the best network of this nature in Latin America. I understand it is run by an independent organization but building a valuable system like this the question of whether it can be integrated and how to a future EEWS pops up naturally.

Gerardo Suárez  
Instituto de Geofísica UNAM

*[Please see the attachment that begins on the next page.]*



36

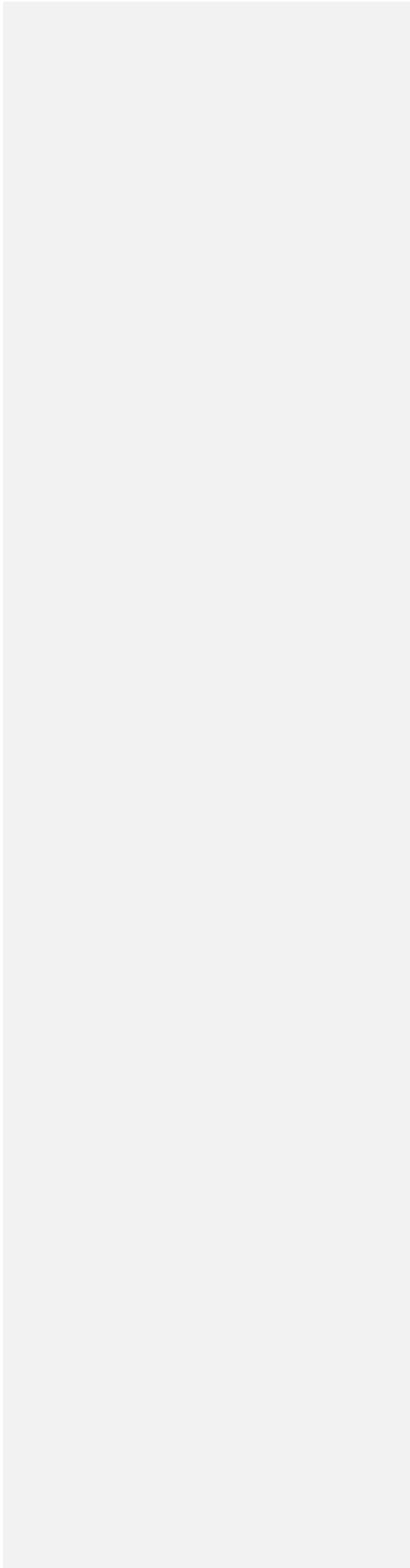
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42 **Abstract**

43 We show using the ASTUTI (Alerta Sismica Temprana Utilizando Teléfonos Inteligentes)  
44 network that smartphones deployed in a fixed network can provide Earthquake Early Warning  
45 performance comparable to scientific-grade systems. Using the phones' accelerometers, we  
46 implement a ground-motion detection and alerting strategy focusing on subduction zone  
47 earthquakes [in Costa Rica](#). Our strategy considers that much of the country's population is  
48 concentrated near the capital San Jose and that nearly the entire population experiences shaking  
49 during earthquakes of  $\sim M_w$  6. Rather than waiting until an earthquake reaches a high intensity,  
50 we evaluate issuing alerts when an acceleration threshold is exceeded at multiple stations. From  
51 more than six months of observations we find data latency is 0.35-0.45 secs. We simulate the  
52 2012 M7.6 Nicoya earthquake using on-phone vibrations and find median first-alert latency of  
53  $\sim$ 9-13 secs after origin time and alert-receipt latency ~~using Amazon Web Services Simple~~  
54 ~~Notification Services~~ of  $\sim$ 4 secs. During our assessment there were 13 earthquakes that caused  
55 felt shaking. From offline reanalysis, we detected and alerted on 5 of these, all of which  
56 produced felt shaking [above ? g](#) in San Jose. The system did not produce any false alerts and  
57 undetected events ([define these events](#)) did not produce wide-spread felt shaking. If we alerted  
58 the entire population for each event, 70% and 15% of the population would receive alerts in time  
59 to undertake drop-cover-hold-on for events inside and outside of the network, respectively.  
60 ~~Complementary-Some~~ population [ranges-centers](#) would receive alerts and not feel shaking. This  
61 strategy may be effective if users are tolerant of feeling no shaking when they receive alerts from  
62 correct detections.

63

64 **Plain Language Summary**

65 We show that a network of smartphones deployed in fixed locations can provide Earthquake  
66 Early Warning performance on par with scientific grade instrumentation. This approach makes  
67 EEW ~~to~~-accessible to resource-limited countries that may ~~have-not been to~~-able to ~~previously~~  
68 benefit from some form of EEW. In particular, combining the fixed-network smartphone sensors  
69 with a strategy targeting low cost ~~of action~~-protective measures such as Drop Cover Hold On  
70 could be very effective.  
71

72 **1 Introduction**

73 Earthquake early warning (EEW) attempts to rapidly detect earthquakes and to alert people and  
 74 systems with enough time for protective actions to be taken before damaging shaking arrives  
 75 [Heaton, 1985]. EEW systems are operational in a handful of regions and they are now rapidly  
 76 being developed and adopted globally [Allen and Melgar, 2019]. As with many techno-scientific  
 77 advances with the potential for large societal impacts, however, the earliest projections of EEW  
 78 system performance and benefits to society have been modified and scaled back as the  
 79 theoretical Earthquake early warning's (EEW) fundamental promise is that earthquakes can be  
 80 rapidly detected so that people and systems can be alerted to take protective action before  
 81 shaking arrives at their location [Heaton, 1985]. In order to maximize warning time and to  
 82 minimize the population not receiving sufficient warning, EEW requires a dense sensor network  
 83 so that earthquakes can be detected closest to ~~wherever they may~~ the point of nucleation. The  
 84 locations where earthquakes nucleate and where people reside, however, might be quite removed  
 85 from one another. This requirement for dense sensor networks, combined with the high cost of  
 86 expensive scientific-grade sensors currently limits EEW systems to wealthy countries [Allen and  
 87 Melgar, 2019] (I don't consider Mexico to be a wealthy country and it built the first public EEWs  
 88 in the world. The financial problem was solved using home made, low cost instruments)].  
 89 Alternatively, a new generation of low-cost accelerometer and geodetic sensors [E. S. Cochran,  
 90 2018] could make EEW generally accessible. In particular, utilizing smartphones, including via  
 91 crowd-sourcing, is a potentially transformative way to provide EEW [Minson et al., 2015;  
 92 Finazzi, 2016; Kong et al., 2016]. Although crowd-sourcing removes sensor cost from EEW  
 93 budgets, it is fundamentally limited, in contrast to fixed networks, by its inability to  
 94 consistently leverage the distance between earthquake source regions and population centers  
 95 (This sentence is awkward. Please rewrite). Smartphones can, of course, also be employed in  
 96 fixed networks. With fixed networks, depending on the location of the event and the position of  
 97 the sensors, all users could potentially get warnings (Figure 1a). By definition, in contrast,  
 98 however, some users in a crowd-sourced EEW system will may never not get warnings (Figure  
 99 1b). Combining the detection advantage with the low capital and communications costs of  
 100 smartphones may make fixed networks ~~an the most~~ optimal low-cost EEW configuration,  
 101 perhaps even comparable to scientific grade systems. To the best of our knowledge, however,  
 102 smartphone fixed network performance has never been examined. In this paper, with a fixed  
 103 network in Costa Rica, we demonstrate that judiciously placed smartphones can provide  
 104 operational EEW for a country's population centers at a cost drastically reduced from, and at a  
 105 performance level equivalent to, scientific grade instrumentation.

106  
 107 As with many techno-scientific advances with the potential for large societal impacts, the earliest  
 108 projections of EEW system performance and its benefits to society have been modified and  
 109 scaled back as EEW's theoretical, empirical, practical, and social limitations are better  
 110 understood [M A Meier, 2017; Sarah E. Minson et al., 2018; S.E. K McBride et al., 2019;  
 111 Minson et al., 2019; Nakayachi et al., 2019; Trugman et al., 2019; Becker et al., 2020; S  
 112 McBride et al., 2020; David J Wald, 2020b], empirical [M A Meier, 2017; Trugman et al., 2019],  
 113 practical [David J. Wald, 2020a], and social [Nakayachi et al., 2019; Becker et al., 2020; S  
 114 McBride et al., 2020] limitations of EEW are elucidated. In parallel with this evolution in  
 115 understanding, EEW instrumentation and network design has also rapidly evolved. Although  
 116 most systems, such as ShakeAlert in the U.S. [Given et al., 2018; Kohler et al., 2020] use  
 117 traditional seismic instrumentation, it has recently become clear that low-cost, consumer-grade

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118 sensors in smartphones including MEMs accelerometers [E Cochran et al., 2009; J Evans et al.,  
119 2014; Kong et al., 2016] and GNSS chips [S.E. Minson et al., 2015] are sensitive enough to be  
120 effectively used in EEW systems, even in crowd-sourced modes [Lawrence et al., 2014; S.E.  
121 Minson et al., 2015; Finazzi, 2016; Kong et al., 2016; Finazzi and Fassò, 2017; Kong et al.,  
122 2018; Kong et al., 2019b; Finazzi, 2020; Kong et al., 2020].

123  
124 Given this rapidly evolving context, it is important to consider that a one-size-fits-all approach to  
125 EEW may not be appropriate, in particular for resource-limited populations. Varied EEW  
126 objectives and performance requirements could likely be achieved with different combinations of  
127 instrumentation, algorithms, alerting strategy, and messaging strategy [Y. Fujinawa et al., 2011].  
128 For instance, EEW objectives for automated industrial systems with large action costs (such as  
129 nuclear power plants) may require highly sensitive instrumentation, high ground-motion  
130 prediction accuracy levels, low false alarm tolerance, and large annual budgetary allocations  
131 [Strauss and Allen, 2016]. Alternatively, if the EEW objective is to promote population  
132 participation in an immediate low-cost-of-action response such as Drop Cover Hold On (DCHO)  
133 [Porter and Jones, 2018], then sensor sensitivity may be lower, ground-motion prediction  
134 accuracy may not be as important, false alarm tolerance may be higher, and budgetary allocation  
135 may be much smaller. Our work here examines this low-budget end of the spectrum.

136  
137 Specifically, we ask: can off-the-shelf smartphones deployed in a fixed network provide EEW in  
138 a reliable and effective enough manner so that significant populations receive warnings in time  
139 for DCHO protective actions to be undertaken? Instead of investing in the institutional overhead  
140 of designing, building, and maintaining our own sensing hardware and communications platform  
141 [Anthony et al., 2019], the smartphone network approach leverages the massive research and  
142 development programs of smartphone providers such that sensors, field computing, and  
143 communications platforms are provided in a single compact unit. Moreover, the constant  
144 competition in this consumer space assures frequent technological updates as instrumentation  
145 becomes obsolete. Although there are now multiple examples of smartphone-based EEW  
146 methodologies and capabilities [S.E. Minson et al., 2015; Finazzi, 2016; Kong et al., 2016;  
147 Finazzi and Fassò, 2017; Kong et al., 2018; Kong et al., 2019b; Kong et al., 2020], to the best of  
148 our knowledge, there have been no evaluations of smartphones used, over extended periods of  
149 time, to provide EEW for a specific region.

150 . To motivate and place in context our work, we review these recent developments, in particular  
151 as they pertain to smartphone-based EEW.

152  
153 Ideally, an EEW system would issue actionable alerts prior to P-(~~compressional~~)-wave arrival,  
154 affording the most time possible for protective actions. Even though a felt P-wave would  
155 hopefully serve as a natural EEW system alerting people to imminent shaking [David J. Wald,  
156 2020a], site and path effects can preclude P-waves being felt ubiquitously, so a minimum  
157 criterion for an EEW system to be effective is for it to deliver alerts prior to the subsequent  
158 arrival of strongest shaking associated with S-waves. Currently, many operational EEW systems  
159 detect earthquakes using some permutation of point-source ~~parameter estimation from the early~~  
160 P-wave. For instance, ShakeAlert uses a small amount of the P-wave ( $>0.2$  secs) from 4 stations  
161 and empirical scaling relations to estimate point source event location and magnitude magnitude  
162 estimation from P-wave information [Chung et al., 2019]. In Japan, P-wave parameter estimation  
163 is augmented with a particle filter technique to combine ground motion observations and

164 ~~information about where shaking has not been observed [Tamaribuchi et al., 2014]. In Mexico,~~  
 165 ~~an event's body wave magnitude,  $m_b$ , is estimated empirically from the rate of seismic energy~~  
 166 ~~released between P and S wave arrivals at a series of free field acceleration stations along the~~  
 167 ~~country's subduction zone coast [Cuéllar et al., 2017]. Point-source-based parameter estimation,~~  
 168 ~~albeit fast at initial detections, exhibits degraded performance when events become large and~~  
 169 ~~magnitude estimates saturate [Hoshihira et al., 2010]. Moreover, theoretical studies have recently~~  
 170 ~~found that the conventional EEW approach of using source parameters to forecast shaking causes~~  
 171 ~~alerts to be too slow for higher levels of shaking [M A Meier, 2017; Sarah E. Minson et al.,~~  
 172 ~~2018; Trugman et al., 2019]; and to mostly produce missed and false alerts due to the intrinsic~~  
 173 ~~factor of two variability of ground motion [Gregor et al., 2014; S.E. Minson et al.,~~  
 174 ~~2019]. Accordingly, EEW practitioners are exploring and/or incorporating alternate~~  
 175 ~~methodologies including line source estimates from strong motion data [Böse et al., 2012],~~  
 176 ~~surface displacements from GNSS data [S. E. Minson et al., 2014], and [Gregor et al., 2014;~~  
 177 ~~Minson et al., 2019]. Accordingly, EEW practitioners are developing ground-motion based~~  
 178 ~~approaches that issue alerts when one or more stations observe shaking above a threshold~~  
 179 ~~[Kodera et al., 2018; Elizabeth S Cochran et al., 2019].~~

180  
 181 ~~In particular, the Propagation of Local Undamped Motion (PLUM) algorithm is a ground-~~  
 182 ~~motion-based approach that was developed as an alternative to source-parameter EEW~~  
 183 ~~methodologies [Kodera et al., 2018; E S Cochran et al., 2019; M A Meier et al., 2020]. PLUM~~  
 184 ~~arose in response to the 2011 M9.0 Tohoku oki earthquake, whose productive aftershock~~  
 185 ~~sequence led to many missed events and false alarms from the point source algorithm when~~  
 186 ~~simultaneous small earthquakes were combined into a single larger event (Hoshihira et al., 2011;~~  
 187 ~~Kodera et al., 2018). Unlike most EEW methods that attempt to characterize source parameters~~  
 188 ~~and forecast expected shaking, PLUM uses shaking to directly forecast shaking. By~~  
 189 ~~instance, using strong ground motion that accrues sometime between P- and S-wave arrivals to directly~~  
 190 ~~forecast ground motion, alerts can be issued for strong shaking as soon as it is observed without~~  
 191 ~~having to wait for the rupture (and earthquake magnitude) to grow in size. PLUM handles inter-~~  
 192 ~~event variability well, reducing false and missed alerts due to inaccuracies in the ground motion~~  
 193 ~~forecast [Kodera et al., 2018; Elizabeth S Cochran et al., 2019].~~

194  
 195 ~~In addition to detection methodology, alerting criteria are also quite variable. ShakeAlert issues~~  
 196 ~~alerts for ground motions predicted to be greater than Modified Mercalli Intensity (MMI) III~~  
 197 ~~[Thakoor et al., 2019]. In Japan, warnings are issued by the JMA (Japan Meteorological Agency)~~  
 198 ~~to seismic hazard blocks [Yukio Fujinawa and Noda, 2013], roughly four per prefecture, when~~  
 199 ~~any evaluation point inside a block is predicted to experience shaking greater than JMA IR~~  
 200 ~~(intensity reading) 4.0 (MMI VI), but only after an IR of 5L (MMI 7) is forecast in at least one~~  
 201 ~~area [Kamigaichi et al., 2009; Kodera et al., 2018]. In Mexico, if rapidly estimated magnitude is~~  
 202 ~~greater than  $m_b$  6, public warnings are issued in Mexico City and other inland locales. Alerting~~  
 203 ~~criteria are further~~  
 204 ~~In addition to detection methodology, EEW alerting criteria are also quite~~  
 205 ~~variable. Alerting criteria are complicated by the simple physical fact that earthquake magnitude~~  
 206 ~~is not static; rather, it increases as an earthquake rupture evolves. The promise of EEW systems~~  
 207 ~~has been amplified by the idea that differences between the ways large and small earthquakes~~  
 208 ~~start- could potentially be used to predict the final size of a growing earthquake, rather than~~  
 209 ~~relying on rapidly updating contemporaneous estimates of source parameters and ground motion~~  
 [Olson and Allen, 2005; Melgar and Hayes, 2019]. However, a growing number of studies have

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210 shown that the early P-wave holds little to no predictive power [*M-A Meier et al., 2017;*  
 211 *Goldberg et al., 2019; Trugman et al., 2019*]. Furthermore, *Minson et al (2018)* recently showed  
 212 that given the velocities at which earthquake ruptures grow and seismic waves propagate,  
 213 together with the exponential decay of ground-motion with distance from a propagating rupture,  
 214 users' warning time for the strongest shaking will often be brief if EEW systems wait to issue  
 215 alerts until earthquakes grow large enough that their forecast shaking is strong. ~~Accordingly, it is~~  
 216 ~~becoming clear that EEW offers the most potential for successful mitigating actions~~ Accordingly,  
 217 it is becoming clear that EEW offers the most potential for successful mitigating actions, such as  
 218 Drop-Cover-Hold-On (DCHO) [*Porter and Jones, 2018*] to be taken if alert thresholds are set at  
 219 lower ground-motion levels than those expected to cause damage [*Sarah E. Minson et al., 2018;*  
 220 *Saunders et al., 2020*].

221  
 222 ~~It is important to consider that while~~ Although an EEW system with low alert thresholds would  
 223 likely produce alerts without ~~subsequent strong~~ shaking in some locations, the number of missed  
 224 alerts would be minimized and the system would have the best chance of being effective ~~when it~~  
 225 ~~mattered most~~ during medium to larger events. For such a strategy to be successful, however,  
 226 users must be more tolerant of receiving unnecessary alerts (the 'Boy Who Cried Wolf'  
 227 phenomena) than is often assumed. Although there are essentially no studies explicitly  
 228 addressing the 'Boy Who Cried Wolf' phenomenon for EEW, other scientific communities with  
 229 more established alerting relationships with users caution against the implicit assumption that  
 230 user false alarm tolerance is low [*Roulston and Smith, 2004*]. In fact, the small amount of  
 231 evidence that exists suggests that EEW users may be ~~surprisingly~~ false-alarm tolerant. *Nakayachi*  
 232 *et al (2019)*, found from surveys associated with alerts from two Japanese events that, indeed, the  
 233 user population was surprisingly tolerant of false alerts. Anecdotal surveys in Mexico found  
 234 similar attitudes [*Allen et al., 2018*].

235  
 236 In this contribution, we implement the ASTUTI (Alerta Sísmica Temprana Utilizando Teléfonos  
 237 Inteligentes; Earthquake Early Warning Utilizing Smartphones) network ~~and we evaluate the~~  
 238 ~~coupling of a fixed smartphone network and ground motion based detection methodology with a~~  
 239 ~~low threshold alert strategy. Smartphone MEMS accelerometers are well suited for ground-~~  
 240 ~~motion based EEW methodologies, having been shown to have noise levels sufficiently low to~~  
 241 ~~permit detection of small (M4-5) earthquakes [*Kong et al., 2016; Finazzi and Fassó, 2017; Kong*~~  
 242 ~~*et al., 2018; Kong et al., 2019a; Kong et al., 2019b; Finazzi, 2020; Kong et al., 2020*]. Although~~  
 243 ~~we introduced the concept of smartphone crowd-sourced EEW [*S.E. Minson et al., 2015*], we~~  
 244 ~~maintain that it is prudent to question crowd-sourcing's operational reliability and overall~~  
 245 ~~practicality. Even though efforts such as The Earthquake Network Project [*Finazzi, 2016*] and~~  
 246 ~~MyShake [*Kong et al., 2016*] report high initial enrollment rates, and even though Google has~~  
 247 ~~recently announced that the Android operating system will have native ability to use~~  
 248 ~~accelerometer data for EEW detection purposes [*Press, 2020*], there is no guarantee that users~~  
 249 ~~will continue to permit their phones to be utilized for extra-personal purposes, especially as~~  
 250 ~~internet privacy advocacy becomes a more penetrative societal issue [*Ketelaar and Van Balen,*~~  
 251 ~~*2018*]. Additionally, depending on citizen subscribers, with their changing preferences and~~  
 252 ~~resources, likely poses a substantial risk in operating a warning system that requires continuous~~  
 253 ~~data from the near source regions of the most damaging expected earthquakes. Finally, as in~~  
 254 ~~Costa Rica. ~~For~~ During six 6 months of continuous operation of the 82 station network, we assess~~  
 255 the coupling of the fixed smartphone network with a ground-motion based detection

256 ~~methodology~~. Please clarify this sentence. In what context do you use the word coupling? We evaluate a low-threshold  
 257 alert strategy that would alert the entire country upon the unambiguous detection of an event  
 258 (Any magnitude?). The spatial distribution of events during the time period appears to be  
 259 representative of the general spatial distribution of Costa Rican seismicity and allows us to  
 260 examine system performance for events that occur within and outside of the network.  
 261 Specifically, for each event, we quantify the percentage of Costa Rican population that would  
 262 have sufficient warning time to undertake DCHO mitigating actions. Finally, all of the raw data  
 263 collected for this project are freely available. As commercial entities with proprietary algorithm  
 264 and data policies are poised to increase by orders of magnitude the number of smartphone EEW-  
 265 capable sensors and to issue public-safety alerts themselves [Stogaitis *et al.*, 2020], we assert that  
 266 it becomes ever more important that the scientific community have access to transparently  
 267 collected and freely available smartphone data.

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Comment [GS3]:

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## 270 2 Materials and Methods

### 271 2.1 System Design

272 Costa Rica's seismic hazard is due primarily, though not exclusively, to earthquakes generated by  
 273 oblique subduction of the Cocos plate below the Caribbean plate at rates up to ~8.5 cm/yr along  
 274 the Middle America Trench (MAT) [DeMets *et al.*, 2010; Protti *et al.*, 2014](Figure 42). In  
 275 contrast to other subduction zones, ~~the portion of the plate interface where~~ MAT earthquakes ~~in~~  
 276 ~~Costa Rica tend to nucleate is overlain by land under the continent and not rather than the~~  
 277 ~~offshore seafloor~~ [Protti *et al.*, 2014]. Since 1853, 8 earthquakes greater than  $M_w$  7.0 have  
 278 occurred either on the northern Nicoya (1853, 1900, 1950, 2012) or southern Osa (1856, 1904,  
 279 1941, 1983) peninsulas [Kobayashi *et al.*, 2014; Protti *et al.*, 2014]. The majority of Costa  
 280 Rica's ~5M population resides in the greater San Jose region, ~60-200 km from the Pacific coast  
 281 and principal seismogenic zone (Figure 42). For greater San Jose as well as Costa Rica's coastal  
 282 regions, 10% expected peak ground acceleration (PGA) exceedance over the next 50 years is  
 283 0.55-0.90g, in the upper ranges of global seismic hazard [Pagani *et al.*, 2020]. Recently, the  $M_w$   
 284 7.6 2012 Nicoya peninsula earthquake ~~caused shaking throughout was felt in~~ the entire country,  
 285 with PGA values as high as 0.5-1.4g and MMI V-VII reported in San Jose (Figure 4b2b), ~~and it~~  
 286 ~~is~~. ~~generally~~ Generally, ~~accepted the~~ earthquakes  $M > 6$  events are felt country-wide.

287  
 288 ~~Given Based on~~ Costa Rica's seismotectonic framework and population distribution we designed  
 289 the ASTUTI ~~network~~ with these principles: (1) Our EEW efforts are focused on warning people,  
 290 not automated systems, and our targeted user response is DCHO. Because the cost of taking  
 291 action for DCHO is so low, this implies that detection and alerting thresholds can be low. (2)  
 292 We prioritize detecting and alerting for MAT earthquakes, the events that have the highest  
 293 probability of affecting the largest percentage of the population, especially ~~in~~ San Jose. Although  
 294 San Jose's central location exposes it to earthquake sources from the entire country, most non-  
 295 MAT sources would be too close to San Jose to permit warnings. (3) If we wait to issue an alert  
 296 until an event has grown large enough to cause shaking damage, then it will most likely be too  
 297 late to issue actionable warnings [Sarah E. Minson *et al.*, 2018; Trugman *et al.*, 2019][Minson *et*  
 298 *al.*, 2018; Trugman *et al.*, 2019]. Accordingly we attempt to issue warnings at the earliest  
 299 detection of events of potential concern. (4) Because of local ground-motion variability ~~S.E.~~

300 | ~~Minson et al., 2019~~[Minson et al., 2019] and because it is unlikely that earthquake ruptures are  
 301 | deterministic [M-A Meier et al., 2017; Goldberg et al., 2019; Trugman et al., 2019] our detection  
 302 | and alerting is entirely non-parametric. In addition to making no attempt to estimate source  
 303 | information (such as location and magnitude), the alert does not include information about  
 304 | predicted ground-motion levels. (5) Based on the previous design principles, Costa Rica's  
 305 | generally small areal extent, and the country-wide shaking from MAT events such as the 2012  
 306 | Nicoya event, we do not attempt to designate intra-country warning precincts; rather we evaluate  
 307 | the scenario where every alert will be issued for the entire country. (6) Because of the previous  
 308 | design priorities, there will likely be a number of alerts issued for smaller events when users will  
 309 | feel no shaking although an event was correctly detected. This will require rapid post-event  
 310 | messaging and constant user interaction and education to remind users that the system performed  
 311 | correctly even if they did not feel shaking [S McBride et al., 2020]. We stress, however, that, as  
 312 | of January 2021, we are not issuing public alerts, aside from those sent to our small group of beta  
 313 | testers. We leave thorough investigation of this topic to a future paper.  
 314 |

## 315 | 2.2 Hardware, Network, and Data Architecture

316 | From September to December 2019 we constructed the ASTUTI network (Figure 1a2a). For the  
 317 | duration of the testing period, the network comprised 82 stations. Closest to the MAT and most  
 318 | of the strongest expected sources, station spacing is ~30 km and it increases to 30-50 km away  
 319 | from the MAT. Higher station density closer to the subduction zone is similar to the ~~SASMEX~~  
 320 | configuration in Mexico [Espinosa-Aranda et al., 2009; Cuéllar et al., 2017; 2018; Suárez et al.,  
 321 | 2018]. Phones are installed on the ground floors of buildings in protective boxes and affixed to  
 322 | floors or walls (Figure 23). Previously, in addition to accelerometer data from smartphones, we  
 323 | have discussed and used GNSS, which typically requires outdoor installation [~~S.E. Minson et al.,~~  
 324 | ~~2015~~][Minson et al., 2015]. We found that many devices overheated, however, so for the initial  
 325 | phase of ASTUTI we made the operational decision to only employ phones installed on interior  
 326 | walls. Accordingly, the only data we use for ASTUTI are from ~~on-board the built-in~~  
 327 | accelerometers. To date, our approach utilizes smartphones with the Android Operating System  
 328 | (~~Android OS~~). Criteria for phone choice included in-country availability and cost (Supp. Mat).  
 329 |

330 | ~~Onboard-Inside~~ the phones, our control and sensing software ~~is called QED (for 'Quick Event~~  
 331 | ~~Detection'; Figure 3; Supp. Mat). QED controls sensor sampling and logging, detects events, and~~  
 332 | ~~prepares data (via formatting, filtering, or on-board data reduction) for use in various~~  
 333 | ~~downstream processing algorithms, and sends data in short, labeled messages. QED harnesses as~~  
 334 | ~~much on-board computation from the smartphone's CPU as possible, thereby eliminating some~~  
 335 | ~~computational load (and potential latency increases) when large numbers of stations are active in~~  
 336 | ~~the network. QED supports either continuous or triggered streaming at sampling rates up to 100~~  
 337 | ~~Hz for accelerometer data- (Figure 4; Supp. Mat). From December 2019 to late August 2020,~~  
 338 | phones streamed at 10 Hz; subsequently we increased streaming rate (~~Streaming in this context is~~  
 339 | ~~equivalent to the sampling rate? Not clear in the discussion)~~ for the entire network to its current  
 340 | rate of 100 Hz. At 100 Hz, data rate for each site is ~70-100 Mb/day.  
 341 |

342 | ~~Data from the phones are streamed using the UDP (User Datagram Protocol) protocol to a cloud-~~  
 343 | ~~based data receiver (Supp. Mat). We prefer UDP to TCP/IP (Transmission Control~~  
 344 | ~~Protocol/Internet Protocol) because it eliminates latency caused by two-way communications~~

Comment [GS4]:

345 ~~regarding packet completeness and, similar to others [JR Evans et al., 2005], we have found~~  
 346 ~~UDP packet loss to be negligible. The UDP receiver passes data to a MQTT (Message Queuing~~  
 347 ~~Telemetry Transport) broker that distributes it to processing and archiving subscribers (Figure~~  
 348 ~~3; As of January 2021 we have collected more than 0.5 Tb of data (please see Acknowledgments,~~  
 349 ~~Samples, and Data section below for information on accessing the archived data). As of January~~  
 350 ~~2021 we have collected more than 0.5 Tb of data. The phones actively maintain NTP (Network~~  
 351 ~~Time Protocol) clock corrections to ensure timing accuracy (Supp. Mat) and they periodically~~  
 352 ~~check in to a settings server comprising a MySQL (My Structured Query Language) database~~  
 353 ~~that provides configurable control parameters. Additionally, a management console provides a~~  
 354 ~~map and table based general monitoring and control interface for the network.~~

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### 355 2.3-PGAN Detection Algorithm and Alerting Strategy

356 Any EEW processing algorithm that operates on acceleration data can subscribe to the MQTT  
 357 broker. Here, we report on a new algorithm, PGAN (for ‘Peak Ground Acceleration with N  
 358 vertices’) that is an adaption of the PLUM method for a network of smartphones. PGAN is  
 359 similar to PLUM in that it is a non-parametric ground-motion based algorithm which does not  
 360 attempt to estimate anything about a detected earthquake’s source, such as magnitude or  
 361 epicentral location. ~~(I believe this discussion is important to include) When PLUM is run as a~~  
 362 ~~component of JMA’s operational EEW system, however, it not only utilizes high quality~~  
 363 ~~seismometers, but data feeds are also monitored by JMA personnel to ensure quality control and~~  
 364 ~~flag noisy data. To compensate for the noisier nature of smartphone MEMS accelerometers~~  
 365 ~~installed in buildings susceptible to anthropogenic noise, as well as for the lack of human~~  
 366 ~~monitoring, PGAN requires multiple neighboring stations to experience anomalous accelerations~~  
 367 ~~in order to trigger an alert. The network is divided into a polygonal mesh with a configurable~~  
 368 ~~number of vertices greater than or equal to three. Here, we report on a quadrilateral (PGAN-~~  
 369 ~~4) Our detection algorithm is a modification of the ground-motion based Propagation of Local~~  
 370 ~~Undamped Motion algorithm [Kodera et al., 2018]. In a polygonal mesh of station locations, we~~  
 371 ~~compensate for the noisier nature of smartphone accelerometers by requiring multiple~~  
 372 ~~neighboring stations to experience anomalous accelerations in order to trigger an alert. Here, we~~  
 373 ~~report on a quadrilateral mesh configuration of adjoining stations because we found an~~  
 374 ~~unacceptably large number of false alerts for triangular configurations.~~

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375 ~~In detail, PGAN-4~~ The station locations are organized into a set of unique, overlapping polygons  
 376 whose sides are all less than a configurable length (currently 40 km). ~~We permit polygons to~~  
 377 ~~overlap.~~ Once per second PGA values are measured at each station and the values are compared  
 378 with a primary threshold (currently 0.6%g). If at any station’s PGA is above the threshold, the  
 379 polygon it belongs to is marked as potentially triggered. The PGA values at each of the  
 380 remaining stations in that polygon are compared with a secondary threshold (currently 0.55%g).  
 381 If PGA values at all four stations in a potentially triggered polygon are above the thresholds, an  
 382 alert is issued. If not, the incoming PGA values continue to be monitored for up to 15 seconds  
 383 (configurable), and the polygon will trigger if the remaining station PGA values rise above the  
 384 secondary PGA threshold. Otherwise, the polygon times out, and no alert is triggered. Once a  
 385 polygon is triggered, an alert is sent out using ~~the a cloud-based notification system (Amazon~~  
 386 ~~Web Services Simple Notification System (AWS SNS)). As of January 2021, alerts are only being~~  
 387 ~~sent to a small number of beta-subscribers in our research team.~~

388

389

### 390 3 Data Quality and System Latency

391 ~~Below, we characterize the system's d~~Data quality and latency was evaluated during a six month  
 392 ~~evaluation~~ period, from December 2019 to June 2020. ~~Generally, d~~During this time, daily system  
 393 up-time (defined as the percentage of each day the entire network was transmitting data) was  
 394 ~~reliably consistengly~~ greater than 95%. After June 2020, as Covid-19 pandemic restrictions  
 395 significantly limited travel in Costa Rica, we were not able to visit stations for routine or time-  
 396 dependent site visits. Accordingly, station up-time and network coverage decreased during this  
 397 ~~period to a ?% level~~. In the Supporting Information section we provide more details about system  
 398 operation.

#### 399 3.1 Data Quality

400 As with any seismic sensor, coupling to the ground and site-specific noise conditions control an  
 401 individual phone's sensitivity. MEMS accelerometers in smart-phones have been shown to have  
 402 the sensitivity to discriminate ground motions caused by earthquakes from background noise ~~f~~  
 403 ~~Evans et al., 2014][Evans et al., 2014]~~, even in crowd-sourcing scenarios where they rest on  
 404 tables [Finazzi, 2016; Kong et al., 2016; Finazzi and Fassò, 2017; Kong et al., 2019b]. In our  
 405 fixed-network approach, phones ~~do not rest on tables nor are they placed in seismically quiet~~  
 406 ~~underground vaults, rather they~~ are affixed to floors, baseboards, and walls in built structures,  
 407 typically homes, schools, fire departments, hospitals, or municipal buildings (Figure 23). As  
 408 such, they record accelerations through the filter of a ~~built~~ structure emplaced in locally varying  
 409 soil and/or rock substrate. This type of installation means that individual stations will record  
 410 earthquakes (Figure 4a5a) as well as ambient non-seismic accelerations ~~from myriad sources~~  
 411 ~~such produced by as~~ road traffic, domestic motion, or thunder [Finazzi, 2020]. In addition to site-  
 412 specific noise, we have found that some phones experience sporadic, unexplained, transient noise  
 413 spikes, either on individual or multiple accelerometer components simultaneously (Figure 4b5b).

414  
 415 ~~To construct the most appropriate data for the PGAN algorithm we~~We combine the raw  
 416 acceleration data from the three orthogonal ~~MEMS~~ accelerometers into a PGA data type (the 'P'  
 417 ~~message~~) sampled and sent at 1 Hz (~~the 'P', for 'PGA', message~~). ~~This P-message data~~ value is  
 418 the vector norm of the individually de-meant 100 Hz acceleration values and has units of  $m/s^2$ ,  
 419 also expressed as a percent of gravitational acceleration, %g. To mitigate ~~the~~ transient noise  
 420 spikes while permitting real seismic accelerations to pass, ~~final formation of the P message~~  
 421 ~~implementwe implement~~ a simple filter that takes the 30<sup>th</sup> percent highest value in a 1 second  
 422 window (3<sup>rd</sup> highest if accelerometer sampling is 10 Hz, 30<sup>th</sup> highest if accelerometer sampling is  
 423 100 Hz; Figure 4b5b) [Kamigaiichi et al., 2009].

424  
 425 Because the ~~PGAN~~ detection algorithm only requires exceedance of ground acceleration above a  
 426 threshold, we characterize ~~each~~ site and ~~network polygon (??)~~ data quality by examining  
 427 histograms of PGA data. Generally, the entire network exhibits mean ~~P~~-values of 0.25%g. Over  
 428 periods of time, varying from minutes to days, individual stations may experience elevated  
 429 deviations from this background behavior (Figure 56) ~~(Please explain units on vertical axis. Not~~  
 430 ~~clear if it is a histogram, as the text says)~~. The causes of these transient periods of elevated  
 431 background noise are varied and likely related to temporally changing site-specific noise such as  
 432 local construction projects, changing traffic patterns, or persistent regional storms.

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433 *3.2 Latency Budget*

434 Understanding and documenting the latency budget, along with assessing alert validity, is a  
 435 critical aspect of EEW system operation. We define ASTUTI system latency,  $\delta t_{\text{latency}}$ , as the sum  
 436 of three components  $\delta t_{\text{latency}} = \delta t_{\text{data}} + \delta t_{\text{detect}} + \delta t_{\text{alert}}$  where  $\delta t_{\text{data}}$  is the time it takes for data to be  
 437 transmitted from the phones and received by the hub,  $\delta t_{\text{detect}}$  is the time it takes for a given  
 438 processing algorithm to detect an event, and  $\delta t_{\text{alert}}$  is the time it takes for a message to be received  
 439 by a user after it has been sent by the processing algorithm. Of these,  $\delta t_{\text{data}}$  and  $\delta t_{\text{alert}}$  are only  
 440 dependent on telecommunications factors and they are independent of the specifics of a given  
 441 earthquake event. For different earthquakes,  $\delta t_{\text{detect}}$  depends on multiple factors including  
 442 magnitude, rise-time, hypocentral location, near-source network geometry, and local site  
 443 acceleration response. These factors will control the time it takes for seismic waves to travel  
 444 from an earthquake to be sensed by a phone, the time it takes for a site to exceed detection  
 445 thresholds, and the time it takes a particular algorithm to issue a detection.

447 From six months of continuously-streamed data, we find that  $\delta t_{\text{data}}$  varies from 0.35 to 0.45  
 448 seconds depending on time of day; peak internet usage (early evening when people arrive home  
 449 from work) also correlates with higher  $\delta t_{\text{data}}$  (Figure 6a7a). For comparison, published  $\delta t_{\text{data}}$  for  
 450 the Italian [Satriano *et al.*, 2011], the Chinese [Zhang *et al.*, 2016], and ShakeAlert [E-S Cochrane  
 451 *et al.*, 2018] [Elizabeth S Cochrane *et al.*, 2018] EEW networks are 0.9, 2, and 1-3 secs,  
 452 respectively. (These are surprisingly low transmission latency times. It would be useful to know  
 453 how they are calculated and show some examples of it).

455 In order to robustly and simultaneously measure the real-time distribution of  $\delta t_{\text{detect}}$  and  $\delta t_{\text{alert}}$ , as  
 456 well as to test the full operation of the ASTUTI EEW system, we used smartphones'  
 457 programmable vibration feature to cause the phones to shake at expected relative arrival times  
 458 for a scenario earthquake. We use the  $M_w$  7.6 2012 Nicoya earthquake as a scenario event and  
 459 program S-wave arrival time for each site calculating hypocentral distance (see methods) using a  
 460 fixed move-out value for  $V_s$  of 3.2 km/s (Figure 6b7b-d) (Figure 7a seems out of place here and is  
 461 not referenced in the text). Each phone vibrated for ~10 seconds. Although they certainly do not  
 462 reproduce the frequency nor amplitude content of real seismic waves interacting with built  
 463 structures, smartphone vibrations are decent-reasonable proxies for local earthquake  
 464 accelerations in that they as they exceed the acceleration threshold values we used for earthquake  
 465 detection at specified times consistent with the earthquake rupture evolution and seismic wave  
 466 propagation (Figure 6e7c). Of course, the model of seismic wave propagation for the event could  
 467 be more sophisticated but the constant  $V_s$  value is sufficient for the purpose of estimating average  
 468 values of  $\delta t_{\text{detect}}$  and  $\delta t_{\text{alert}}$ . Note here we use S-wave arrival times to be conservative, but it is  
 469 possible that trigger thresholds could be exceeded in the P-wave (see System Performance  
 470 below).

471  
 472 Over a period of 3 days, we repeated the test, vibrating the phones on the Nicoya peninsula three  
 473 times an hour at the M7.6 scenario relative times, resulting in a total of 216 simulations (Figure  
 474 6d7d). For this scenario, we find  $\delta t_{\text{detect}}$  for the PGAN-4 method has a mean value of ~12-13  
 475 seconds. We measured  $\delta t_{\text{alert}}$  by sending text message alerts to a set of 15 people with cell phones  
 476 in and around the greater San Jose region. We find mean value for  $\delta t_{\text{alert}}$  is ~4 seconds (Figure  
 477 6e7d). Because this is from a small number of phones that were relatively close together, it is not

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478 clear how representative this metric is of latencies that would occur when sending many  
 479 thousands of alerts across a wider geographic region. For comparison we note, that in a recent  
 480 test of ShakeAlert in San Diego county using the U.S. federal Wireless Emergency Alert (WEA)  
 481 system [[Sarah E. Minson et al., 2020](#)][[Minson et al., 2020](#)], median  $\delta t_{\text{alert}}$  was  $\sim 13$  seconds. To  
 482 the best of our knowledge no other smartphone EEW projects have reported alerting latency data  
 483 [[Finazzi, 2016](#); [Kong et al., 2016](#); [Finazzi and Fassò, 2017](#); [Kong et al., 2018](#); [Kong et al.,](#)  
 484 [2019a](#); [Kong et al., 2019b](#); [Finazzi, 2020](#); [Kong et al., 2020](#)].  
 485

## 486 4 System Performance

### 487 4.1 Event Accuracy and Timeliness

488 We evaluate ASTUTI's performance in both detecting earthquakes and delivering alerts to  
 489 people who will potentially experience shaking for a given earthquake, based on off-line  
 490 playback from Dec 2019 to June 2020. ~~Because of Covid-19 related travel restrictions limiting  
 491 field work and equipment maintenance, and because of algorithm tuning and development of our  
 492 alert messaging algorithm, the entire system was not operational until August 2020.~~

493 We use Did You Feel It (DYFI) data to provide both a self-consistent metric and the most natural  
 494 ground-truth for whether people in a given region experienced shaking [[Atkinson and Wald,](#)  
 495 [2007](#)]. ~~If, alternatively, our objective were to evaluate performance for non-human 'users' such  
 496 as critical facilities, lifeline, and/or structures, then DYFI reports would be less appropriate than  
 497 instrumental strong motion measurements such as ShakeMaps [[David J Wald, 2000](#)].~~

499 During the period of assessment, ~~ASTUTI using the PGAN-4 algorithm we~~ detected 5 of the 13  
 500 events ( $M_w$  4-5.3) that were accompanied by DYFI reports of shaking ~~somewhere~~ in Costa Rica  
 501 (Figure [78](#), Figure S1, Table S1). The majority of these events (9 of 13) occurred outside of the  
 502 network, either offshore in the MAT or ~~in within neighboring in~~ Panama's international borders  
 503 (Figure S1). Four of the detected events had thrust mechanisms and one had a strike-slip  
 504 mechanism and ~~they~~ ranged in magnitude from from  $M_w$  4.8 ~~to  $M_w$  to~~ 5.3. Only one of the  
 505 detected events (2020-03-07,  $M_w$  5.2) had an epicenter entirely within the network (Figure ~~748d,~~  
 506 [8d](#)). ~~Generally, the ASTUTI detected events were accompanied by stronger shaking that was felt  
 507 by much larger percentages of the population as defined by interpolated DYFI reports.~~ The  
 508 detected events had a median MMI of 4.3 and a max MMI of 6 with  $\sim 17\%$  to  $73\%$  (41%  
 509 median) of the population experiencing felt shaking of at least MMI 2-3. In contrast, the non-  
 510 detected events had median MMI of 2.9 and max MMI of 3.8 with 0 to 19% (0.002% median) of  
 511 the population experiencing felt shaking of at least MMI 2-3.  
 512

513 For each detected event, in addition to the associated DYFI data we plot the estimated position of  
 514 the S-wave front (assumed to be the front of peak shaking) at the time that the alert was issued  
 515 (solid magenta circles in Figure [78](#)). Detection times,  $\delta t_{\text{detect}}$  ~~for PGAN-4~~ ranged from 11-30  
 516 secs (median 22 secs, Table 1, Figure [87](#)). ~~(In the Supplemental Material section we present  
 517 results using the PGAN-3 configuration exhibiting median  $\delta t_{\text{detect}}$  of 16.8 seconds).~~ The 11-  
 518 30 second range of detection times compares well with other scientific grade systems constructed  
 519 and operated in similar seismic hazard settings, such as Cascadia and Mexico. For Cascadia, in  
 520 2019-2020 Shake Alert issued 4 alerts with  $\delta t_{\text{detect}}$  timing ranging from 8.3-13.9 seconds

521 [ShakeAlertEventPage, 2020]. Published ~~examples from the SASMEX system in Mexico show~~  
 522  $\delta t_{\text{detect}}$  values ~~from Mexico~~ for 3 events ~~ranging range~~ from 12-18 seconds [Cuéllar *et al.*, 2018]  
 523 and, although not explicitly presented, we estimate from data presented in ~~the~~ recent-SASMEX  
 524 performance summary [Cuéllar *et al.*, 2017] that minimum  $\delta t_{\text{detect}}$  is  $\sim 8$  seconds. However, in  
 525 the Mexican system there is essentially no data transmission nor data distribution delay, as the  
 526 alert is defined locally by the sensing stations and sent to users via radio links.

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527  
 528 For each detected event we plot record sections ~~for each station's recorded~~ of P messages and of  
 529 the expected range of P- and S-wave velocities of 6 to 3 km/s, respectively (Figure 89). We find  
 530 that 2 of the 5 events (2019-12-08 and 2020-03-07, Figure 8a9a,b and Figure 8e9g,h) triggered  
 531 on the P wave, ~~which can be seen as shown~~ by the magenta vertical line ~~occurring~~ on the left side  
 532 of the record section plots. For each of these events, all four of the triggering-PGAN stations  
 533 exceeded the threshold (0.6 g for the first and 0.55 %g for subsequent stations; shown in dashed  
 534 green lines on the right panel ~~in-of~~ each figure). The three remaining events triggered with the S-  
 535 wave. For all ~~of the~~ events, ~~except aside from that on the one on~~ 2020-03-07,  $\sim 10$  seconds  
 536 elapsed between the first and fourth station detections, surpassing the detection threshold. For the  
 537 2020-03-07 event, only 2 seconds elapsed, because the 2020-03-07 event was entirely within the  
 538 network ~~and so spreading seismic waves from the sub-surface event arrived much more~~  
 539 ~~synchronously at the sensors than for events with offshore hypocenters.~~

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#### 541 4.2 Shaking

542 In addition to reporting the detection times we also assess the likely warning outcomes for  
 543 population percentages ~~for the case where the warning precinct for a detected event is the entire~~  
 544 ~~Costa Rican population (Figure 7). We use the WorldPop 1 km gridded database~~ [Tatem, 2017]  
 545 ~~for our population data: the case where the warning precinct for a detected event is the entire~~  
 546 ~~Costa Rican population (Figure 8).~~ We interpolate the spatial distribution of DYFI reports to  
 547 define an area of felt shaking, reasoning that it is likely ground-shaking occurred between sites of  
 548 reported shaking [Kodera *et al.*, 2018; ~~Elizabeth S Cochran~~ *et al.*, 2019]. Although DYFI data  
 549 contains uncertainty from humans' perceptual subjectivity [Goltz *et al.*, 2020], we minimize  
 550 this by applying a binary "shaking" or "no-shaking" classification to DYFI reports. We contend  
 551 this is a conservative approach in that, for a given earthquake, the area of felt shaking  
 552 represented by DFYI data is an under- rather than over-estimation: it seems more likely that  
 553 people who felt shaking would not report it rather than that people would report felt shaking  
 554 when none was felt. We categorize outcomes as True Positive-Shaking (TP-S): the system  
 555 correctly detects an event and a user receives an alert prior to felt shaking; True Positive-No  
 556 Shaking (TP-NS): the system correctly detects an event, a user receives an alert but does not feel  
 557 shaking; No Alert (NA): the system correctly detects an event, but a user does not receive an  
 558 alert prior to shaking; False Alert (FA): the system incorrectly detects an event and sends an  
 559 alert, no shaking occurs anywhere; and Missed Alert (MA): a felt event occurs but the system  
 560 does not issue an alert. For each of these scenarios, to permit conservative evaluation of when  
 561 users could expect to receive warning messages, in the plots we add to  $\delta t_{\text{detect}}$  5 seconds,  $\sim 1$   
 562 second more than the median value of  $\delta t_{\text{alert}}$  from the vibration tests (see section 3.2).

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563 ~~For PGAN 4~~ Over the 6 months of evaluation, 15.5-71.0% (median 38.5%) of the population  
 564 would have received TP-S outcomes, 27.1-83% (median 52.8 %) would have TP-NS outcomes,  
 565

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566 and 4.9-8.8% (median 2.5%) would have NA outcomes (Table 1, ~~in Table S2 we display metrics~~  
 567 ~~for PGAN 3 configuration~~). Outcomes can be grouped further according to onshore (in-  
 568 network) and offshore (out-of-network) source categories. For the onshore events (12/8/19,  
 569 3/7/20, and 3/13/20) more time after warning is received is generally available because  $\delta t_{\text{detect}}$  is,  
 570 on average, 10 seconds faster than the two offshore events (1/21/20a,b) and because the events  
 571 on either the Nicoya or Osa peninsulas are far enough away from population concentrations  
 572 (Figure 7a8a,d,e). Interestingly, however, the two off-shore events yield the largest TP-S  
 573 outcomes because Costa Rica's population is concentrated in and around the capital city, San  
 574 Jose, near the center of the country and more than 100 km from these events (Table 1).  
 575

#### 576 4.3 Missed & False Alerts

577 In order to analyze a self-consistent dataset of felt shaking, we further classify MAs to be events  
 578 for which the system did not produce an alert but that produced enough shaking for people to file  
 579 DYFI reports. This is a minimum number as there may be other, smaller events that were felt but  
 580 for which no DYFI reports were filed. During the period of evaluation there were 8 such MAs  
 581 with a range of  $M_w$  4-5 (Table S1, Figure S1). Six of these MAs were located well outside of the  
 582 network (Figure S1). Of the remaining two, only one (12-31-2020,  $M_w$  4) was located entirely  
 583 within the network, the other (4-17-2020,  $M_w$  4) was located on the Pacific coast at the edge of  
 584 the network. For these events, less than 0.5% of the population reported shaking and the reported  
 585 shaking never exceeded MMI III (Figure S1).  
 586

587 False alerts (FAs) can occur because of system glitches or non-seismic local or regional  
 588 accelerations. From our time period of analysis we identified two types of ~~system glitch~~ sources  
 589 that affected three or more neighboring stations nearly simultaneously. The first is unexpected  
 590 phone vibration because of user error when programming the vibration tests. The second is  
 591 regional power-grid instability causing AC electricity fluctuations that, in turn, cause multiple  
 592 phones to unexpectedly vibrate simultaneously when they re-start as the regional AC power  
 593 cycles off and on. Once they are identified, these are relatively straightforward FA sources to  
 594 mitigate in real-time by better accounting of programmed vibrations and by identifying and  
 595 excluding sites experiencing regional power cycling. ~~Real, non-earthquake accelerations from~~  
 596 ~~sources such as vehicles passing on the street, thunder shaking a structure, or people moving~~  
 597 ~~furniture near the phone installation are more difficult to mitigate. As the number of vertices in~~  
 598 ~~the PGAN triggering criteria is increased, the probability of regionally simultaneous non-seismic~~  
 599 ~~accelerations decreases. In our early tests we found that it was necessary to include more than~~  
 600 ~~three stations in the PGAN detection algorithm to mitigate the effect of these types of unwanted~~  
 601 ~~accelerations on alert generation. (In my opinion this discussion is unnecessary because you~~  
 602 ~~require at least four sensors in the polygon to report exceedance of the PGA. Traffic, human~~  
 603 ~~motion or the other effects you mention would not take place simultaneously)~~  
 604

605 ~~In our analysis of six months of ASTUTI data using PGAN 4~~ During the evaluation period, once  
 606 system glitches were identified and noisy stations removed, we eliminated all FAs yielding a FA  
 607 rate of 0%. This is similar ~~performance~~ to the result from offline runs ~~of the PLUM algorithm~~ on  
 608 West Coast ~~U.S. data that had a 0% FA rate of the ground-motion based algorithm from which~~  
 609 ~~ours is derived~~ (Cochran, *et al.* 2019). For further comparison, published FA rates for science-  
 610 grade EEW networks range from 2.5% in Taiwan [Xu *et al.*, 2017] to 8% for Shake Alert [Kohler

Comment [GS9]:

611 *et al.*, 2020]. The Japanese system reported many FAs after the 2011 *M* 9 Tohoku earthquake but  
 612 to the best our knowledge has not published a FA rate [Hoshiya, 2014]. For smartphone EEW,  
 613 the Earthquake Network Project sets detection thresholds so that they do not exceed 1 FA per  
 614 year, per country [Finazzi, 2020]. This approach is a way to balance the often competing  
 615 objectives of minimizing both MAs and FAs. The MyShake project has not reported FA rate,  
 616 stating that “the false positive events are not so well quantified, because we do not have a large  
 617 number of false positive samples from the system” [Kong *et al.*, 2020].  
 618

## 619 5 Discussion

620 Our investigation demonstrates for the first time that off-the-shelf smartphones deployed in a  
 621 fixed network can reliably and effectively provide operational EEW for a country with Costa  
 622 Rica’s size and population. Performance metrics for the ASTUTI network compare favorably  
 623 with scientific-grade EEW, at least during the six months of evaluation. Data latency is lower  
 624 than some science grade networks and detection latency is similar albeit slightly higher due to  
 625 the ground-motion based detection method triggering usually on S-waves. The spatial  
 626 distribution of earthquakes that occurred during the evaluation period appears to be  
 627 representative of the spatial distribution of the earthquakes expected to do the most damage in  
 628 Costa Rica (with the exception of the southern Caribbean region) and so it is likely that similar  
 629 ASTUTI performance could be expected over longer time intervals.

630  
 631 There is a marked cost differential between the fixed network smartphone approach and science  
 632 grade networks. Capital costs for the entire ASTUTI network were ~USD 22,000 and annual  
 633 operating and maintenance costs are ~USD 20,000 (not 1-2 full time equivalent salaries). For  
 634 comparison, ShakeAlert capital costs are close to USD 100 million and annual operating costs  
 635 are ~ USD 39 million [Given *et al.*, 2018]. When normalized by population for Costa Rica (~5  
 636 M) and combined populations of California, Oregon, and Washington (~52 M), ASTUTI’s  
 637 capital costs are 450 times less and annual operating costs are 192 times less than Shake Alert’s.  
 638 When normalized by area for the same regions (~51,100 km<sup>2</sup> and ~863,000 km<sup>2</sup>, respectively),  
 639 ASTUTI’s capital costs are 279 times less and annual operating costs are 119 times less than  
 640 Shake Alert’s. Despite this roughly two order of magnitude difference in capital and annual  
 641 costs, we find, generally, that performance metrics for our fixed network smartphone EEW  
 642 approach compare favorably with scientific grade EEW, at least during our six months of  
 643 evaluation. Data latency,  $\delta t_{data}$ , for the ASTUTI network is lower than some science grade  
 644 networks. Alerting latency,  $\delta t_{alert}$ , is the same regardless of the type of sensing network.  
 645 Detection latency,  $\delta t_{detect}$ , is similar, albeit slightly higher. Increasing station density by ~ 30%  
 646 such that four stations would cover the area that three stations currently cover (at a cost of less  
 647 than ~USD 10,000) would/could probably decrease  $\delta t_{detect}$  by 3-5 seconds. Furthermore, the spatial  
 648 distribution of earthquakes that occurred during the first six months of ASTUTI network  
 649 operation appears to be representative of the spatial distribution of the earthquakes expected to  
 650 do the most damage in Costa Rica (with the exception of the southern Caribbean region) and so  
 651 it is likely that similar ASTUTI performance could be expected over longer time intervals.

652  
 653 Faster detection times for EEW systems using higher grade sensors is due primarily to P-wave  
 654 based detection [Cuéllar *et al.*, 2018; Chung *et al.*, 2019]. We identify P waves in the  
 655 acceleration records of individual stations for all 5 detected events (Figure 8) and, notably, two

656 ~~of the five events were detected as the result of triggers in the P waves. Accordingly, although~~  
 657 ~~ground-motion based algorithms may be more likely to trigger on S waves for moderate sized~~  
 658 ~~events, there is no *a priori* requirement of this. Faster warnings could occur for larger events,~~  
 659 ~~such as the  $M_w$  5.2 2020-03-07 event where P arrivals triggered the event and are seen in the~~  
 660 ~~records of many stations (Figure 8d). Additionally, as future generation MEMs accelerometers~~  
 661 ~~become more sensitive, detection on the P wave could become more likely, although trigger~~  
 662 ~~thresholds have to be carefully considered given the noise levels at station locations. A better~~  
 663 ~~understanding of the relationship between expected P wave amplitudes and their spatial~~  
 664 ~~coherence for a variety of events would allow for improved predictions of the performance of~~  
 665 ~~accelerometer based EEW.~~

667 The ASTUTI strategy of warning the entire country upon detection of any event would permit  
 668 warnings to reach large percentages of the population (15-70% TP outcome) for the earthquakes  
 669 that “mattered”, that is, earthquakes that caused enough shaking for people to submit DYPF  
 670 reports. The detected events had median and maximum MMI levels of 4.3 and 6 with ~17% to  
 671 73% (41% median) of the population experiencing felt shaking. In contrast, the non-detected  
 672 events had median and max MMI levels of 2.9 and 3.8 with 0 to 19% (0.002% median) of the  
 673 population experiencing felt shaking. For smaller and sufficiently out-of-network events that did  
 674 not trigger alerts, less than 0.5% of the population reported shaking and the reported shaking  
 675 never exceeded MMI H3. These population outcomes are a function of the spatial distribution of  
 676 event locations, station density, and population spatial distribution. For example, for events that  
 677 nucleate on the Nicoya or Osa peninsulas, the detecting stations are far enough away from  
 678 population concentrations to afford more people time to receive warnings and undertake  
 679 mitigating actions. Consistent with the results of the vibration test for the 2012 M7.6 Nicoya  
 680 earthquake scenario, this suggests that the approach could work well for the case of subduction  
 681 zone events of significant size (>M7). ~~For events in the central Pacific coastal embayment~~  
 682 ~~between the Nicoya and Osa peninsulas, however, the events (2020-01-21a,b) occur farther away~~  
 683 ~~from the coastal sensors, and closer to higher population centers, leading to shorter detection and~~  
 684 ~~warning times. This is an unavoidable consequence of Costa Rica’s geography and is common to~~  
 685 ~~any EEW system reliant on land based sensors.~~

687 ~~For outcomes where an alert is received and no shaking experienced, we believe it is important~~  
 688 ~~to differentiate between event based FA and shaking based TP-NS outcomes. The~~  
 689 ~~ASTUTI/ASTUTI’s 0% FA rate is lower than all other EEW systems’ reported FA rates aside~~  
 690 ~~from retrospective testing of ground-motion based detection with West Coast U.S. data [E-S~~  
 691 ~~Cochran et al., 2019][Elizabeth S Cochran et al., 2019]. Our examination was only over six~~  
 692 ~~months, however, and so more operational time is required for a longer-term, more~~  
 693 ~~representative FA rate. As expected, a low detection threshold criterion combined with a~~  
 694 ~~country-wide alerting region also leads to significant population percentages (20-80%) with~~  
 695 ~~where an alert was received and no was shaking experienced (TP-NS outcomes-). TP-NS~~  
 696 ~~percentage could be reduced by modifying our alerting philosophy to include shaking estimation~~  
 697 ~~from source-parameter characterization; however, we suggest that the benefit of attempting a~~  
 698 ~~more refined warning may be outweighed by the added uncertainty associated with EEW~~  
 699 ~~parameter estimation and ground-motion prediction [Sarah E. Minson et al., 2018; S.E. Minson~~  
 700 ~~et al., 2019][Minson et al., 2018; Minson et al., 2019]. It is not clear what penalty there may be~~  
 701 ~~in terms of user engagement if an EEW system provides alerts without felt shaking, especially if~~

702 users were to receive a rapid follow-on message stating that an event had been correctly detected  
 703 although they did not feel shaking. In fact, the general population may appreciate receiving an  
 704 alert whenever an earthquake occurred, even if no one felt it [Nakayachi *et al.*, 2019]. Indeed, we  
 705 are at the very early stages of studying the nuanced relationships between EEW system  
 706 performance and human sentiment. Given this evolving understanding, a clear high-priority must  
 707 be EEW pre- and post-event education and messaging [S McBride *et al.*, 2020].  
 708

709 Given the concentration of Costa Rica's population in San Jose (> 90%), it is useful to ~~further~~  
 710 ~~consider outcomes~~ assess DCHO feasibility for ~~its~~ the greater metro region. S-wave arrival in San  
 711 Jose for the three onshore or near-shore events (2019-12-08, 2020-03-07, 2020-03-13) was  
 712 between ~30 and 50 seconds after median receipt of alert time ( $\delta t_{\text{detect}} + 5$  seconds) (Figure  
 713 ~~7a8a~~, d, e). For the offshore events (2020-01-21a, 2020-01-21b) the S-wave arrival was much  
 714 sooner, a few seconds after median receipt of alert time (Figure ~~7b8b~~, c). Thus, for three of the  
 715 five detected events, our results suggest that DCHO for a large percentage of San Jose residents  
 716 is an achievable objective. For the two offshore events DCHO may not be a widely achievable  
 717 objective. Because there is a paucity of studies that have directly evaluated DCHO, however,  
 718 there is some uncertainty in assessing EEW users' likely response. For instance, for the  
 719 HayWired exercise [Porter and Jones, 2018], DCHO took 5-15 seconds and for a New Zealand  
 720 ShakeOut exercise DCHO timing varied from 10-30 seconds (64% of participants took less than  
 721 10 seconds and 34% took 11-30 secs) [S K McBride *et al.*, 2019]. For two actual shaking events,  
 722 however, Nakayashi *et al.* (2019) report that taking no action was the most common outcome for  
 723 people who received EEW alerts. This uncertainty highlights the often counter-intuitive nature of  
 724 humans' interaction with automated warning systems and stresses the importance of further  
 725 mitigation education and social science studies targeting DCHO activities.  
 726

727 ~~Our investigation began by questioning whether off-the-shelf smartphones deployed in a fixed~~  
 728 ~~network could provide EEW in a reliable and effective enough manner so that significant~~  
 729 ~~populations could receive warnings in time for DCHO protective actions to be undertaken. The~~  
 730 ~~answer is unequivocally affirmative. Faster detection times for EEW systems using higher-grade~~  
 731 ~~sensors is due primarily to P-wave based detection [Cuéllar *et al.*, 2018; Chung *et al.*, 2019]. We~~  
 732 ~~identify P-waves in the acceleration records of individual stations for all 5 detected events~~  
 733 ~~(Figure 9) and, notably, two of the five events were detected as the result of triggers in the P-~~  
 734 ~~waves. Accordingly, although ground-motion-based algorithms may be more likely to trigger on~~  
 735 ~~S-waves for moderate-sized events, there is no *a priori* requirement of this. Faster warnings~~  
 736 ~~could occur for larger events, such as the  $M_w$  5.2 2020-03-07 event where P arrivals triggered the~~  
 737 ~~event and are seen in the records of many stations (Figure 9d). Additionally, as future generation~~  
 738 ~~MEMs accelerometers become more sensitive, detection on the P wave could become more~~  
 739 ~~likely, although trigger thresholds have to be carefully considered given the noise levels at~~  
 740 ~~station locations. A better understanding of the relationship between expected P wave amplitudes~~  
 741 ~~and their spatial coherence for a variety of events would allow for improved predictions of the~~  
 742 ~~performance of accelerometer-based EEW.~~  
 743

744 As pointed out by others [David J Wald, 2020b], ~~however~~, updating construction codes and  
 745 practices as well as promoting more widespread and thorough acceptance of simple mitigation  
 746 actions such as DCHO deserve to be considered ~~as~~ top priorities for earthquake-related resource  
 747 allocation. The ASTUTI fixed network EEW approach ~~we describe here~~ is so inexpensive,

748 | however, that it need not interfere with other budgetary priorities for ~~resource-limited~~ countries  
749 | that do not have the resources to maintain advanced seismic networks but that could benefit from  
750 | some form of EEW. Myriad implementation scenarios ~~can be envisioned ranging from stand-~~  
751 | ~~alone smartphone networks to augmentations of preexisting scientific-grade instrumentation.~~  
752 | ~~Basic stand-up for a network such as ASTUTI requires purchasing and installation of phones and~~  
753 | ~~SIM cards and spinning up instances of the processing algorithms on a cloud-based server.~~  
754 | ~~Similarly, the approach is so inexpensive that the benefits of a fixed-network-smartphone~~  
755 | ~~approach (station coverage for expected earthquake source regions and network reliability)~~  
756 | ~~outweigh the even lower costs associated with crowd-sourcing, could be envisioned ranging~~  
757 | ~~from stand-alone smartphone networks to hybrid augmentations of preexisting scientific-grade~~  
758 | ~~instrumentation. Furthermore, the approach is so inexpensive that the sensing and detection~~  
759 | ~~benefits of judiciously placed sensors outweighs the even lower costs associated with crowd-~~  
760 | ~~sourcing. We note that in no way do we discount crowd-sourcing's value, especially the potential~~  
761 | ~~contribution to ground-motion and structural seismic response studies of massive numbers of~~  
762 | ~~crowd-source acceleration records [Kong et al., 2018].~~

763 | Moreover, as digital home assistants become more ubiquitous, crowd-sourcing efforts (users  
764 | permitting access to their personal devices) will move to fixed network modes. Especially for  
765 | resource-limited populations, however, the day is still years away when user-contributed data  
766 | will be spatially complete enough and temporally reliable enough to provide robust EEW  
767 | capability for a larger population. With smartphone-based EEW in a nascent phase of rapid  
768 | growth, we stress the need for continued scientific investigation and validation of these  
769 | methodologies by the global community. Especially given the need for transparency when issues  
770 | of public safety are concerned, we encourage open-access data policies for smartphone-based  
771 | EEW systems.

## 772 | **6 Conclusions**

773 | We have demonstrated that a fixed-network of smartphones using a ground-motion-based  
774 | detection algorithm can provide operational EEW at a cost that is generally 1-2 orders of  
775 | magnitude less than scientific-grade networks. The ASTUTI combination of detection capability  
776 | with a country-wide alerting strategy demonstrates an effective EEW strategy for the common  
777 | earthquake hazard setting where large population centers are close enough to be affected by large  
778 | earthquakes but far enough away from them so that alerts may arrive in time to permit simple  
779 | protective actions such as DCHO. More social science research is needed to ascertain whether  
780 | populations will be tolerant of a system that correctly detects events and that sends alerts to users  
781 | who do not necessarily feel shaking.

782 |

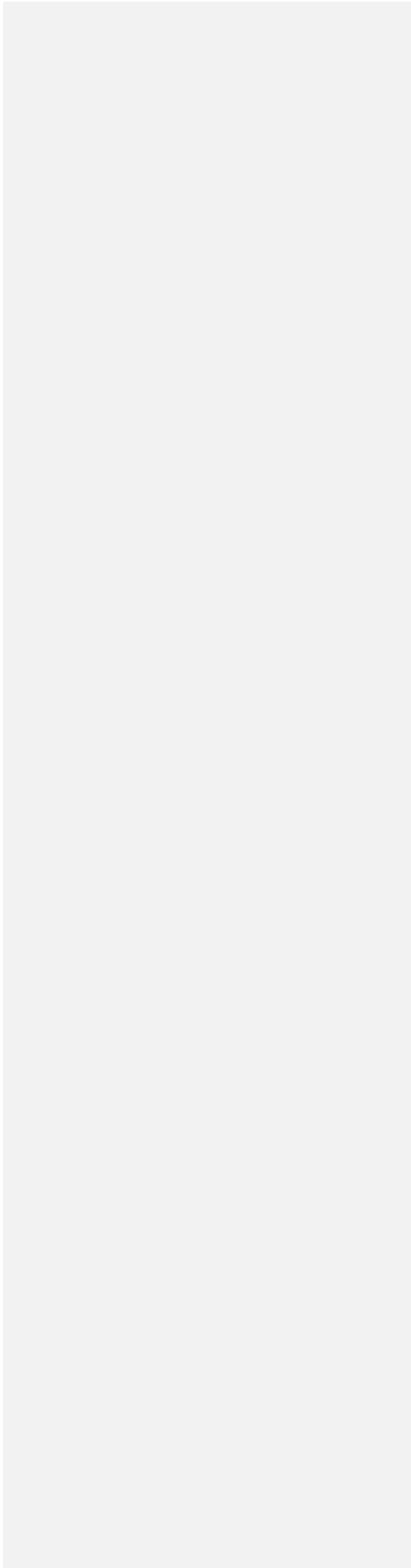
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788 | Data supporting the conclusions in this paper may be obtained at:  
789 | <http://ds.iris.edu/ds/nodes/dmc/forms/assembled-data/>. Assembled data set number: 21-004;  
790 | short name: ASTUTI.

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793



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795 **References**

796

797 Allen, R. M., E. S. Cochran, T. J. Huggins, S. Miles, and D. Otegui (2018), Lessons from Mexico's earthquake early  
798 warning system, *Eos, Earth and Space Science News*, 99.799 Allen, R. M., and D. Melgar (2019), Earthquake early warning: Advances, scientific challenges, and societal needs,  
800 *Annual Review of Earth and Planetary Sciences*, 47, 361-388.801 [Anthony, R. E., A. T. Ringler, D. C. Wilson, and E. Wolin \(2019\), Do low-cost seismographs perform well enough  
802 for your network? An overview of laboratory tests and field observations of the OSOP Raspberry Shake 4D,  
803 \*Seismological Research Letters\*, 90\(1\), 219-228.](#)804 Atkinson, G. M., and D. J. Wald (2007), "Did You Feel It?" intensity data: A surprisingly good measure of  
805 earthquake ground motion, *Seismological Research Letters*, 78(3), 362-368.806 Becker, J. S., S. H. Potter, L. J. Vinnell, K. Nakayachi, S. K. McBride, and D. M. Johnston (2020), Earthquake early  
807 warning in Aotearoa New Zealand: a survey of public perspectives to guide warning system development,  
808 *Humanities and Social Sciences Communications*, 7(1), 1-12.809 [Böse, M., T. H. Heaton, and E. Hauksson \(2012\), Real-time finite fault rupture detector \(FinDer\) for large  
810 earthquakes, \*Geophysical Journal International\*, 191\(2\), 803-812.](#)811 Chung, A. I., I. Henson, and R. M. Allen (2019), Optimizing earthquake early warning performance: ElarmS - 3,  
812 *Seismological Research Letters*, 90(2A), 727-743.813 Cochran, E., [J. Lawrence, C. Christensen, and R. Jakka \(2009\), The Quake-Catcher Network: Citizen Science  
814 Expanding Seismic Horizons, \*Seismological Research Letters\*, 80\(1\), 26. S. \(2018\), To catch a quake, \*Nat Commun\*,  
815 9\(1\), 2508, doi:10.1038/s41467-018-04790-9.](#)816 Cochran, E. S., J. Bunn, S. E. Minson, A. S. Baltay, D. L. Kilb, Y. Kodera, and M. Hoshiba (2019), Event Detection  
817 Performance of the PLUM Earthquake Early Warning Algorithm in Southern California, *Bulletin of the  
818 Seismological Society of America*, 109(4), 1524-1541.819 Cochran, E. S., M. D. Kohler, D. D. Given, S. Guiwits, J. Andrews, M. A. Meier, M. Ahmad, I. Henson, R. Hartog,  
820 and D. Smith (2018), Earthquake early warning ShakeAlert system: Testing and certification platform,  
821 *Seismological Research Letters*, 89(1), 108-117.822 Cuéllar, A., G. Suárez, and J. M. Espinosa - Aranda (2017), Performance Evaluation of the Earthquake Detection  
823 and Classification Algorithm 2(tS-tP) of the Seismic Alert System of Mexico (SASMEX), *Bulletin of the  
824 Seismological Society of America*, 107(3), 1451-1463, doi:10.1785/0120150330.825 Cuéllar, A., G. Suárez, and J. M. Espinosa - Aranda (2018), A Fast Earthquake Early Warning Algorithm Based on  
826 the First 3 s of the P - Wave Coda, *Bulletin of the Seismological Society of America*, 108(4), 2068-2079,  
827 doi:10.1785/0120180079.828 DeMets, C., R. G. Gordon, and D. F. Argus (2010), Geologically current plate motions, *Geophysical Journal  
829 International*, 181(1), 1-80.830 Espinosa-Aranda, J. M., A. Cuellar, A. Garcia, G. Ibarrola, R. Islas, S. Maldonado, and F. H. Rodriguez (2009),  
831 Evolution of the Mexican seismic alert system (SASMEX), *Seismological Research Letters*, 80(5), 694-706 %@  
832 1938-2057.833 Evans, J., R. M. Allen, A. Chung, E. Cochran, R. Guy, M. Hellweg, and J. Lawrence (2014), Performance of several  
834 low - cost accelerometers, *Seismological Research Letters*, 85(1), 147-158.835 [Evans, J. R., R. H. Hamstra Jr., C. Kündig, P. Camina, and J. A. Rogers \(2005\), TREMOR: A wireless MEMS  
836 accelerograph for dense arrays, \*Earthquake Spectra\*, 21\(1\), 91-124.](#)837 Finazzi, F. (2016), The earthquake network project: Toward a crowdsourced smartphone - based earthquake early  
838 warning system, *Bulletin of the Seismological Society of America*, 106(3), 1088-1099.839 Finazzi, F. (2020), The Earthquake Network Project: A Platform for Earthquake Early Warning, Rapid Impact  
840 Assessment, and Search and Rescue, *Frontiers in Earth Science*, 8, 243.841 Finazzi, F., and A. Fassò (2017), A statistical approach to crowdsourced smartphone-based earthquake early  
842 warning systems, *Stochastic Environmental Research and Risk Assessment*, 31(7), 1649-1658.843 [Fujinawa, Y., and Y. Noda \(2013\), Japan's earthquake early warning system on 11 March 2011: performance,  
844 shortcomings, and changes, \*Earthquake Spectra\*, 29\(S1\), S341-S368.](#)

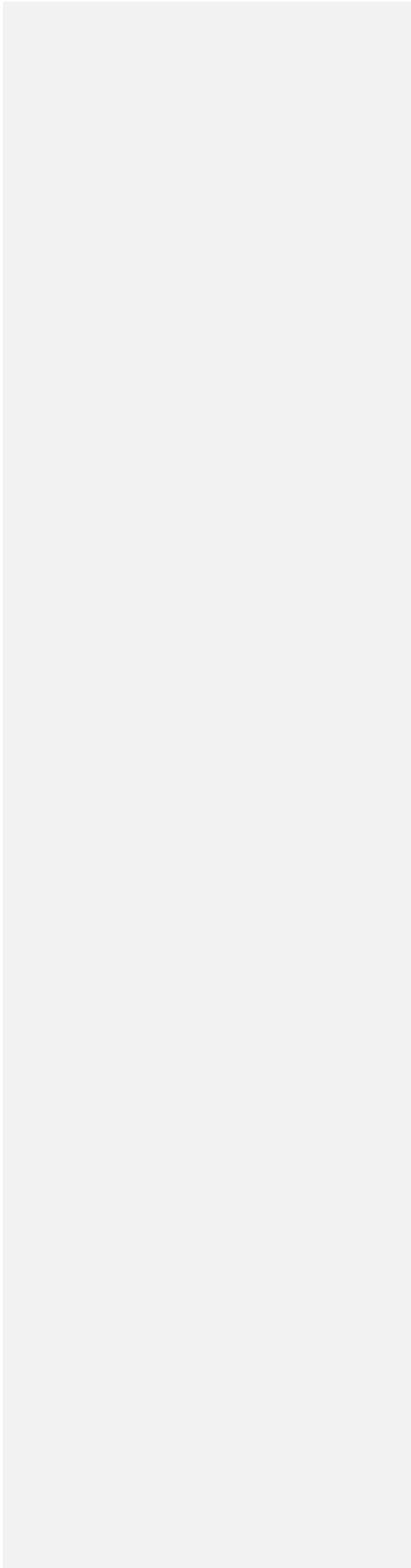
- 845 [Fujinawa, Y., K. Takahashi, Y. Noda, H. Itaka, and S. Yazaki \(2011\), Remote detection of the electric field change](#)  
846 [induced at the seismic wave front from the start of fault rupturing, \*Int. J. Geophys.\*, 752193,](#)  
847 [doi:10.1155/2011/752193.](#)
- 848 Given, D. D., R. M. Allen, A. S. Baltay, P. Bodin, E. S. Cochran, K. Creager, R. M. de Groot, L. S. Gee, E.  
849 Hauksson, and T. H. Heaton (2018), Revised technical implementation plan for the ShakeAlert system—An  
850 earthquake early warning system for the West Coast of the United States *Rep. 2331-1258*, US Geological Survey.
- 851 Goldberg, D., D. Melgar, and Y. Bock (2019), Seismogeodetic P - wave amplitude: No evidence for strong  
852 determinism, *Geophysical research letters*, 46(20), 11118-11126.
- 853 Goltz, J. D., H. Park, G. Nakano, and K. Yamori (2020), Earthquake ground motion and human behavior: Using  
854 DYFI data to assess behavioral response to earthquakes, *Earthquake Spectra*, 8755293019899958.
- 855 Gregor, N., N. A. Abrahamson, G. M. Atkinson, D. M. Boore, Y. Bozorgnia, K. W. Campbell, B. S.-J. Chiou, I.  
856 Idriss, R. Kamai, and E. Seyhan (2014), Comparison of NGA-West2 GMPEs, *Earthquake Spectra*, 30(3), 1179-  
857 1197.
- 858 Heaton, T. (1985), A Model for a Seismic Computerized Alert Network, *Science*, 228, 987-990.
- 859 Hoshiha, M. (2014), Review of the nationwide earthquake early warning in Japan during its first five years, in  
860 *Earthquake Hazard, Risk and Disasters*, edited, pp. 505-529, Elsevier.
- 861 Hoshiha, M., K. Ohtake, K. Iwakiri, T. Aketagawa, H. Nakamura, and S. Yamamoto (2010), How precisely can we  
862 anticipate seismic intensities? A study of uncertainty of anticipated seismic intensities for the Earthquake Early  
863 Warning method in Japan, *Earth, planets and space*, 62(8), 611-620.
- 864 Kamigaichi, O., M. Saito, K. Doi, T. Matsumori, S. y. Tsukada, K. Takeda, T. Shimoyama, K. Nakamura, M.  
865 Kiyomoto, and Y. Watanabe (2009), Earthquake early warning in Japan: Warning the general public and future  
866 prospects, *Seismological Research Letters*, 80(5), 717-726.
- 867 [Ketelaar, P. E., and M. Van Balen \(2018\), The smartphone as your follower: The role of smartphone literacy in the](#)  
868 [relation between privacy concerns, attitude and behaviour towards phone embedded tracking, \*Computers in Human\*](#)  
869 [Behavior](#), 78, 174-182.
- 870 Kobayashi, D., P. LaFemina, H. Geirsson, E. Chichaco, A. A. Abrego, H. Mora, and E. Camacho (2014),  
871 Kinematics of the western Caribbean: Collision of the Cocos Ridge and upper plate deformation, *Geochemistry,*  
872 *Geophysics, Geosystems*, 15, 1671-1683.
- 873 Kodera, Y., Y. Yamada, K. Hirano, K. Tamaribuchi, S. Adachi, N. Hayashimoto, M. Morimoto, M. Nakamura, and  
874 M. Hoshiha (2018), The propagation of local undamped motion (PLUM) method: A simple and robust seismic  
875 wavefield estimation approach for earthquake early warning, *Bulletin of the Seismological Society of America*,  
876 108(2), 983-1003.
- 877 Kohler, M. D., D. E. Smith, J. Andrews, A. I. Chung, R. Hartog, I. Henson, D. D. Given, R. de Groot, and S.  
878 Guiwits (2020), Earthquake Early Warning ShakeAlert 2.0: Public Rollout, *Seismological Research Letters*, 91(3),  
879 1763-1775.
- 880 Kong, Q., R. M. Allen, M. D. Kohler, T. H. Heaton, and J. Bunn (2018), Structural Health Monitoring of Buildings  
881 Using Smartphone Sensors, *Seismological Research Letters*, 89(2A), 594-602, doi:10.1785/0220170111.
- 882 Kong, Q., R. M. Allen, L. Schreier, and Y.-W. Kwon (2016), MyShake: A smartphone seismic network for  
883 earthquake early warning and beyond, *Science advances*, 2(2), e1501055.
- 884 Kong, Q., A. Inbal, R. M. Allen, Q. Lv, and A. Puder (2019a), Machine learning aspects of the MyShake global  
885 smartphone seismic network, *Seismological Research Letters*, 90(2A), 546-552.
- 886 Kong, Q., R. Martin - Short, and R. M. Allen (2020), Toward Global Earthquake Early Warning with the MyShake  
887 Smartphone Seismic Network, Part 1: Simulation Platform and Detection Algorithm, *Seismological Research*  
888 *Letters*.
- 889 Kong, Q., S. Patel, A. Inbal, and R. M. Allen (2019b), Assessing the sensitivity and accuracy of the MyShake  
890 smartphone seismic network to detect and characterize earthquakes, *Seismological Research Letters*, 90(5), 1937-  
891 1949.
- 892 [Lawrence, J. F., E. S. Cochran, A. Chung, A. Kaiser, C. M. Christensen, R. Allen, J. W. Baker, and e. al. \(2014\),](#)  
893 [Rapid earthquake characterization using MEMS accelerometers and volunteer hosts following the M 7.2 Darfield,](#)  
894 [New Zealand, earthquake, \*Bull. Seism. Soc. Am.\*, 104\(1\), 184-192.](#)
- 895 McBride, S., A. Bostrom, J. Sutton, R. M. de Groot, A. S. Baltay, B. Terbush, P. Bodin, M. Dixon, E. Holland, and  
896 R. Arba (2020), Developing post-alert messaging for ShakeAlert, the earthquake early warning system for the west  
897 coast of the United States of America, *International Journal of Disaster Risk Reduction*, 101713.
- 898 McBride, S. K., J. S. Becker, and D. M. Johnston (2019), Exploring the barriers for people taking protective actions  
899 during the 2012 and 2015 New Zealand ShakeOut drills, *International journal of disaster risk reduction*, 37,  
900 101150.

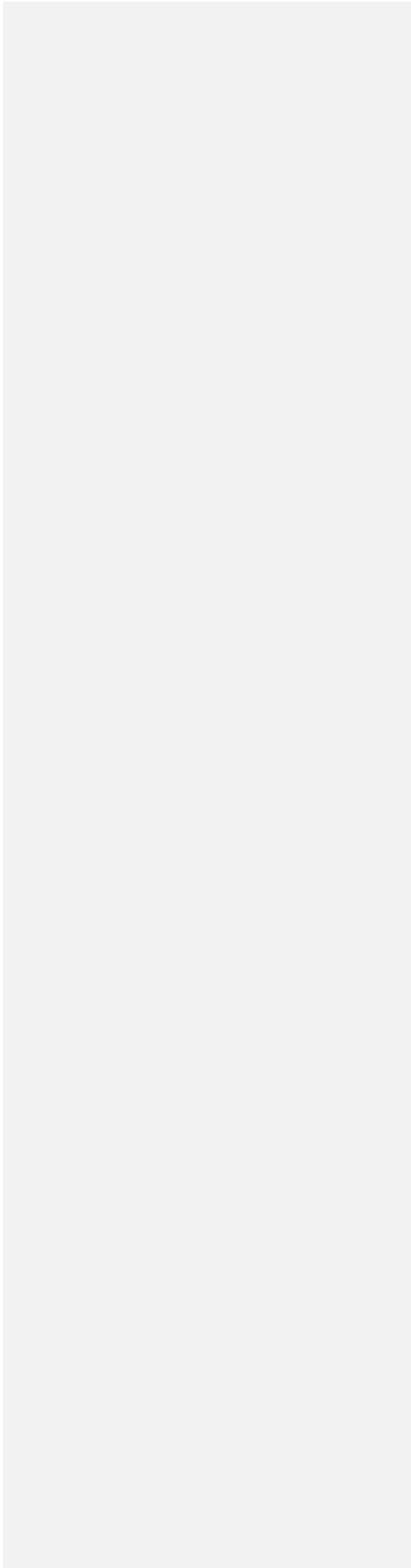
- 901 Meier, M.-A., J. Ampuero, and T. H. Heaton (2017), The hidden simplicity of subduction megathrust earthquakes,  
 902 *Science*, 357(6357), 1277-1281.
- 903 Meier, M. A. (2017), How “good” are real - time ground motion predictions from earthquake early warning  
 904 systems?, *Journal of Geophysical Research: Solid Earth*, 122(7), 5561-5577.
- 905 ~~Meier, M. A., Y. Kodera, M. Böse, A. Chung, M. Hoshiba, E. Cochran, S. Minson, E. Hauksson, and T. Heaton~~  
 906 ~~(2020), How Often Can Earthquake Early Warning Systems Alert Sites With High -Intensity Ground Motion?,~~  
 907 ~~*Journal of Geophysical Research: Solid Earth*, 125(2), e2019JB017718.~~
- 908 Melgar, D., and G. P. Hayes (2019), Characterizing large earthquakes before rupture is complete, *Science Advances*,  
 909 5(5), eaav2032.
- 910 Minson, S. E., A. S. Baltay, E. S. Cochran, T. C. Hanks, M. T. Page, S. K. McBride, K. R. Milner, and M.-A. Meier  
 911 (2019), The Limits of earthquake early Warning Accuracy and Best Alerting strategy, *Scientific reports*, 9(1), 2478  
 912 %@ 2045-2322.
- 913 Minson, S. E., B. A. Brooks, C. L. Glennie, J. R. Murray, J. O. Langbein, S. E. Owen, R. A. Iannucci, and D. L.  
 914 Hauser (2015), Crowd-Sourced Global Earthquake Early Warning, *Science Advances*.
- 915 Minson, S. E., M.-A. Meier, A. S. Baltay, T. C. Hanks, and E. S. Cochran (2018), The limits of earthquake early  
 916 warning: Timeliness of ground motion estimates, *Science advances*, 4(3), eaaq0504 %@ 2375-2548.
- 917 Minson, S. E., J. R. Murray, J. O. Langbein, and J. S. Gombert (2014), Real-time inversions for finite fault slip  
 918 models and rupture geometry based on high-rate GPS data, *Journal of Geophysical Research: Solid Earth*, 119(4),  
 919 doi:10.1002/2013JB01062.
- 920 ~~Minson, S. E., J. K. Saunders, J. J. Bunn, E. S. Cochran, A. S. Baltay, D. L. Kilb, M. Hoshiba, and Y. Kodera~~  
 921 ~~(2020), Real-Time Performance of the PLUM Earthquake Early Warning Method during the 2019 M 6.4 and 7.1~~  
 922 ~~Ridgecrest, California, Earthquakes, *Bulletin of the Seismological Society of America*, 110(4), 1887-1903,~~  
 923 ~~doi:10.1785/0120200021.~~
- 924 Nakayachi, K., J. S. Becker, S. H. Potter, and M. Dixon (2019), Residents’ Reactions to Earthquake Early Warnings  
 925 in Japan, *Risk Analysis*, 0(0 %@ 0272-4332), doi:10.1111/risa.13306 %U  
 926 <https://onlinelibrary.wiley.com/doi/abs/10.1111/risa.13306>.
- 927 Olson, E. L., and R. M. Allen (2005), The deterministic nature of earthquake rupture, *Nature*, 438(7065), 212-215.
- 928 Pagani, M., J. Garcia-Pelaez, R. Gee, K. Johnson, V. Poggi, V. Silva, M. Simionato, R. Styron, D. Viganò, and L.  
 929 Danciu (2020), The 2018 version of the global earthquake model: hazard component, *Earthquake Spectra*,  
 930 36(1 suppl), 226-251.
- 931 Porter, K., and J. Jones (2018), How many injuries can be avoided in the HayWired Scenario through earthquake  
 932 early warning and drop, cover, and hold on, *The HayWired Earthquake Scenario: US Geological Survey Scientific*  
 933 *Investigations Report—2017-5013*. Reston, VA: USGS, 401-429.
- 934 ~~Press, A. (2020), [https://www.usnews.com/news/best-states/california/articles/2020-08-11/california-quake-alerts-](https://www.usnews.com/news/best-states/california/articles/2020-08-11/california-quake-alerts-to-be-standard-on-android-phones)~~  
 935 ~~[to-be-standard-on-android-phones](https://www.usnews.com/news/best-states/california/articles/2020-08-11/california-quake-alerts-to-be-standard-on-android-phones), edited.~~
- 936 Protti, M., V. González, A. V. Newman, T. H. Dixon, S. Y. Schwartz, J. S. Marshall, L. Feng, J. I. Walter, R.  
 937 Malservisi, and S. E. Owen (2014), Nicoya earthquake rupture anticipated by geodetic measurement of the locked  
 938 plate interface, *Nature Geoscience*, 7(2), 117-121.
- 939 Roulston, M. S., and L. A. Smith (2004), The boy who cried wolf revisited: The impact of false alarm intolerance on  
 940 cost-loss scenarios, *Weather and Forecasting*, 19(2), 391-397.
- 941 Satriano, C., L. Elia, C. Martino, M. Lancieri, A. Zollo, and G. Iannaccone (2011), PRESTo, the earthquake early  
 942 warning system for southern Italy: Concepts, capabilities and future perspectives, *Soil Dynamics and Earthquake*  
 943 *Engineering*, 31(2), 137-153.
- 944 Saunders, J. K., B. T. Aagaard, A. S. Baltay, and S. E. Minson (2020), Optimizing Earthquake Early Warning Alert  
 945 Distance Strategies Using the July 2019 Mw 6.4 and Mw 7.1 Ridgecrest, California, Earthquakes, *Bulletin of the*  
 946 *Seismological Society of America*, 110(4), 1872-1886, doi:10.1785/0120200022.
- 947 ShakeAlertEventPage (2020), edited, <https://earthquake.usgs.gov/earthquakes/eventpage/us60007f42/executive>;  
 948 <https://earthquake.usgs.gov/earthquakes/eventpage/uw61562126/executive>;  
 949 <https://earthquake.usgs.gov/earthquakes/eventpage/nc73355700/executive>;  
 950 <https://earthquake.usgs.gov/earthquakes/eventpage/nc73344735/executive>.
- 951 Stogaitis, M., et al. (2020), Earthquakes at Google, presented at 2020 Fall Meeting, American Geophysical Union,  
 952 Abstract S044-08.
- 953 ~~Strauss, J. A., and R. M. Allen (2016), Benefits and costs of earthquake early warning, *Seismological Research*~~  
 954 ~~*Letters*, 87(3), 765-772.~~

- 955 Suárez, G., J. M. Espinosa - Aranda, A. Cuéllar, G. Ibarrola, A. García, M. Zavala, S. Maldonado, and R. Islas  
956 (2018), A dedicated seismic early warning network: The Mexican Seismic Alert System (SASMEX), *Seismological*  
957 *Research Letters*, 89(2A), 382-391 %@ 0895-0695.
- 958 ~~Tamaribuchi, K., M. Yamada, and S. Wu (2014), A new approach to identify multiple concurrent events for~~  
959 ~~improvement of earthquake early warning, *Zisin*, 67(2).~~
- 960 Tatem, A. J. (2017), WorldPop, open data for spatial demography, *Scientific data*, 4(1), 1-4.
- 961 ~~Thakoor, K., J. Andrews, E. Hauksson, and T. Heaton (2019), From earthquake source parameters to ground—~~  
962 ~~motion warnings near you: The ShakeAlert earthquake information to ground—motion (eqInfo2GM) method,~~  
963 ~~*Seismological Research Letters*, 90(3), 1243-1257.~~
- 964 Trugman, D. T., M. T. Page, S. E. Minson, and E. S. Cochran (2019), Peak ground displacement saturates exactly  
965 when expected: Implications for earthquake early warning, *Journal of Geophysical Research: Solid Earth*, 124(5),  
966 4642-4653.
- 967 Wald, D. J. (2000), "*ShakeMaps*": *Instant Maps of Earthquake Shaking*, US Department of the Interior, US  
968 Geological Survey.
- 969 Wald, D. J. (2020a), Practical limitations of earthquake early warning, *Earthquake Spectra*, 36(3), 1412-1447,  
970 doi:10.1177/8755293020911388.
- 971 Wald, D. J. (2020b), Practical limitations of earthquake early warning, *Earthquake Spectra*, 8755293020911388.
- 972 Xu, Y., J. P. Wang, Y.-M. Wu, and H. Kuo-Chen (2017), Reliability assessment on earthquake early warning: A  
973 case study from Taiwan, *Soil Dynamics and Earthquake Engineering*, 92, 397-407,  
974 doi:10.1016/j.soildyn.2016.10.015.
- 975 Zhang, H., X. Jin, Y. Wei, J. Li, L. Kang, S. Wang, L. Huang, and P. Yu (2016), An earthquake early warning  
976 system in Fujian, China, *Bulletin of the Seismological Society of America*, 106(2), 755-765.
- 977 ~~1.~~

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1. Figure 1.





981 **Figure 1.** Cartoon depicting fixed-network (a) and crowd-sourced (b) EEW warning times for  
982 two population centers for a hypothetical earthquake (red star). The red circle indicates the time  
983 when the event is detected. With a fixed-network, the event is detected by nearby sensors located  
984 in zones of sparse population and both population centers could receive warning. With crowd-  
985 sourcing, the first population center's smartphones are used as the detectors and so it does not  
986 receive a warning. Moreover, the warning time for the 2<sup>nd</sup> population center is reduced from the  
987 what the fixed-network could have provided.

988

989 **Figure 2.** Costa Rica seismotectonic framework, population distribution, ASTUTI network, and  
990 the M7.6 2012 Nicoya peninsula earthquake. Basemap, WorldMap 1 km gridded population  
991 (copper colormap) and shaded topography and bathymetry (blue and gray colormaps). (a)  
992 Colored circles, earthquakes in the NEIC catalog from August 15, 2000 to June 3, 2020, colored  
993 by magnitude. White triangles, ASTUTI smartphone locations. MAT, Middle America Trench.  
994 SJ, San Jose city. (b) Beachball, focal mechanism of the 2012 Nicoya earthquake. Colored grid  
995 cells, Did You Feel It (DYFI) Modified Mercalli Intensity (MMI) values from the 2012 Nicoya  
996 earthquake.

997

998 **Figure 23.** A typical ASTUTI station. (a) Photo showing the smartphone enclosure affixed to a  
999 wall. (b) Photo inside the enclosure of the smartphone screen showing the QED software  
1000 application display.

1001

1002 **Figure 34.** Schematic diagram of the ASTUTI approach. The general flow of alerting data and  
1003 information is clockwise starts in the Device Networks box and end in the Alert System box.  
1004 **QED, 'Quick Event Detection' smartphone software.** NTP, network time protocol. UDP, user  
1005 datagram protocol. MQTT, message queuing telemetry transport. API, application programming  
1006 interface. MySQL, My structure queried language.

1007

1008 **Figure 45.** Examples of raw accelerometer and PGA data types from one station. Black, raw  
1009 accelerometer data. V, vertical axis; H1, horizontal axis; H2, horizontal axis. Blue, P, processed  
1010 PGA data message. (a) Example of how the PGA data type preserves the signal of 3 small  
1011 earthquakes at 0, ~75 and ~210 seconds. (b) Example of how the PGA data type suppresses  
1012 noise spikes at ~0 seconds.

1013

1014 **Figure 56.** Histogram of PGA values for one station for one day (gray) and one hour (red). The  
1015 current triggering threshold is 0.6%g.

1016

1017 **Figure 67.** ASTUTI system latency and M7.6 Nicoya scenario vibration test. **(a)** Hourly  
 1018 distribution of data latency,  $\delta t_{\text{data}}$ , for the 6 month observation period. Each column represents  
 1019 the median  $\delta t_{\text{data}}$  probability distribution function for an hour of the day. **(b)** Location map (as in  
 1020 Figure 4a2a) showing the M7.6 Nicoya scenario vibration test. Red star, epicenter. Magenta  
 1021 circle, S-wave position at the time,  $\delta t_{\text{detect}}$ , the time the event was detected. Green triangles, the  
 1022 stations that triggered the alert. **(c)** P-message record section from the vibration test. Stations are  
 1023 ordered with distance from the epicenter. Hypocentral distance shown in blue text. **(d)**  
 1024 Histograms of time to detection,  $\delta t_{\text{detect}}$  (top) and time to receipt of alert,  $\delta t_{\text{alert}}$ , (bottom) for all  
 1025 216 vibration tests.

1026 **Figure 78.** ASTUTI results from the 5 detected earthquakes ~~for the PGAN-4 algorithm~~. Each  
 1027 plot has the same symbology and nomenclature. Title with white box, origin time and magnitude  
 1028 of the event from OVSICORI. Magenta-box title,  $\delta t_{\text{detect}}$ . Magenta solid circle, estimated position  
 1029 of S-wave (3.2km/s) at time  $\delta t_{\text{detect}}$ . Magenta dashed circle, estimated position of S-wave 20  
 1030 seconds after time  $\delta t_{\text{detect}}$ . 20 seconds represents  $\sim 5$  seconds for alerting time,  $\delta t_{\text{alert}}$  (see Figure  
 1031 5d6d) and 15 seconds for protective action such as Drop-Cover-Hold-On (DCHO). Green filled  
 1032 triangles, four stations that triggered the alert. White solid triangles, active stations at time of  
 1033 event. White empty triangles, inactive stations at time of event. Colored grid squares, DYFI  
 1034 cells. Hot population colormap, population % True-positive (TP) outcomes. Gray population  
 1035 colormap, % No-alert (NA). Copper population colormap, % True-positive no-shaking (TP-NS).  
 1036 **(a)** 12 December 2019  $M_w$  4.8. **(b)** 21 January 2020a  $M_w$  5.3. **(c)** 21 January 2020b  $M_w$  5.2. **(d)** 7  
 1037 March 2020  $M_w$  5.2. **(e)** 13 March 2020  $M_w$  5.0.

1038 **Figure 8-9. PGA.** P-message, record section for the 5 detected earthquakes ~~for the PGAN-4~~  
 1039 ~~algorithm~~. Each plot has the same symbology, nomenclature, and scale. The left column shows  
 1040 all stations for a particular event, the right column shows a zoomed image of only the triggering  
 1041 stations. Stations are ordered with distance from the epicenter. Hypocentral distance shown in  
 1042 blue along the vertical axis. Triggering stations, thick black. Other stations, black. Triggering  
 1043 time, magenta line. Pink shading, ranges of 6 and 3 km/s for P- and S-wave arrival times,  
 1044 respectively. **(a,b)** 12 December 2019  $M_w$  4.8. **(c,d)** 21 January 2020a  $M_w$  5.3. **(e,f)** 21 January  
 1045 2020b  $M_w$  5.2. **(g,h)** 7 March 2020  $M_w$  5.2. **(i)** 13 March 2020  $M_w$  5.0.

1046 **Table 1.** Alerting outcomes for each detected event ~~for PGAN-4 algorithms~~. See also Figures 7  
 1047 and 8. Time to detection,  $\delta t_{\text{detect}}$ . TP%, percent of population with True-positive outcomes. TP-  
 1048 NS%, percent of population with True-positive no-shaking outcomes. NA%, percent of  
 1049 population with No-Alert outcomes.

1050  
 1051

### Reviewer #3

The authors are developing a low-cost observation network system over a wide area for earthquake early warnings using smartphones. In addition, the delay time is considered using a large number of seismic records, and it can be evaluated as a highly practical study.

Although it is not related to the essence, I think that it will be a better paper if the following points are corrected, so please consider the correction.

p3 LINE 56

It says "70% and 15% of the population", but according to the description in the text at LINE 494, it would be "15-75%".

p.6 LINE 183

It seems to be an explanation of Figure 2a. Therefore, it is desirable to change the reference from Figure 2 to Figure 2a.

p.10 LINE 363

The word "from" continues twice in a sentence.

p.10 LINE 364

Figure 8d is duplicated, one should be deleted.

p.11 LINE 405

DFYI is misspelled. DYFI is correct.

p.16 LINE 633

The quotes contain unknown characters such as "% @".

p.17 LINE 698,702

The quote contains unknown characters such as "% @".

p.17 LINE 708

The quote contains unknown characters such as "0% @" and "% U".

Also, the volume number is missing. (Risk Analysis, Vol. 39, No. 8, 2019.)

p.18 LINE 736

The quote contains unknown characters such as "% @".

p.18 LINE 751

The unknown character string "1." should be deleted.

Figure 3

Labels (a),(b) are not written on the figure.

Figure 6

The 1-hour histogram is 0.4 to 0.6 more frequently on the horizontal axis than the 1-day histogram. Is the vertical axis common? The vertical axis label should be shown.

Figure 7C

It seems that the terminal with a distance of 30.3km is shaking earlier than the terminal with a distance of 9.9km from the epicenter. I think it is necessary to explain this.

Also, it looks like there are three terminals that aren't shaking at all. Isn't it necessary to explain this?

Table 1

I think that TP-NS is correct when combined with the description in the text.