

SIMULTANEOUS, IN SITU MEASUREMENTS OF OH AND HO<sub>2</sub> IN THE STRATOSPHERE

R. M. Stimpfle, P. O. Wennberg, L. B. Lapson and J. G. Anderson

Department of Earth and Planetary Sciences and Department of Chemistry  
Harvard University, Cambridge, MA 02138

**Abstract.** Stratospheric OH and HO<sub>2</sub> radical densities have been measured between 36 and 23 km using a balloon-borne, in situ instrument launched from Palestine, TX on August 25, 1989. OH is detected using the laser-induced fluorescence technique (LIF) employing a Cu-vapor-laser pumped dye laser coupled with an enclosed-flow detection chamber. HO<sub>2</sub> is detected nearly simultaneously by adding NO to the sample flow to convert ambient HO<sub>2</sub> to OH. Observed OH and HO<sub>2</sub> densities ranged from  $8.0 \pm 2.8 \times 10^6$  and  $1.4 \pm 0.5 \times 10^7$  molec cm<sup>-3</sup>, respectively, at 36 km, to  $1.4 \pm 0.5 \times 10^6$  and  $3.0 \pm 1.0 \times 10^6$  at 23 km, where the uncertainty is  $\pm 1\sigma$ . The HO<sub>2</sub> density exhibits a maximum in the 34-30 km region of  $1.7 \pm 0.6 \times 10^7$ . The data were obtained over a solar zenith angle variation of 51° at 36 km to 61° at 23 km. O<sub>3</sub> and H<sub>2</sub>O densities also were measured simultaneously with separate instruments.

## Introduction

The HO<sub>x</sub> family of reactants is dominated by the OH and HO<sub>2</sub> species in the stratosphere. The photochemical reactions that control HO<sub>x</sub> and the role of HO<sub>x</sub> species in the production and loss of stratospheric ozone are critical to predictions of temporal trends in ozone column abundances (see for example WMO [1986]). In the upper stratosphere, OH and HO<sub>2</sub> engage in odd-oxygen recombination by direct reaction with O<sub>3</sub> and O. In the middle and lower stratosphere, OH and HO<sub>2</sub> affect O<sub>3</sub> indirectly, through their interaction with the NO<sub>x</sub> and Cl<sub>x</sub> families of reactants. The response of OH to increasing Cl<sub>x</sub>, methane and NO<sub>x</sub> in the atmosphere is crucial in determining the effect of these compounds on stratospheric ozone.

A rudimentary method of testing stratospheric HO<sub>x</sub> photochemistry has been by direct comparison of individually measured OH and HO<sub>2</sub> densities (and H<sub>2</sub>O<sub>2</sub> because it is a key indicator of HO<sub>2</sub>) with model calculations. In such comparisons, HO<sub>2</sub> measurements have shown the most striking disparity between measurement and theory. The pioneering measurements of stratospheric HO<sub>2</sub> by cryogenic matrix sampling, followed by ground-based laboratory ESR detection [Mihelcic et al., 1978; Helten et al., 1984], are higher by an order of magnitude or more than model predictions. An in situ technique based upon OH resonance lamp-induced fluorescence detection of OH formed from the conversion of HO<sub>2</sub> to OH by NO addition [Anderson et al., 1981] gave results in better agreement with model predictions but still larger than expected. A ground based, mm-wave spectroscopy, column HO<sub>2</sub> measurement [de Zafra et al., 1984] that is most sensitive

to density variations in the 38-65 km interval did not support the HO<sub>2</sub> densities reported below 37 km by the aforementioned investigators. The disagreement among these HO<sub>2</sub> measurements led to questions concerning the validity of each measurement. The discrepancy between model calculations and the measurements below 38 km, if they are indeed valid, would imply serious errors in our understanding of stratospheric photochemistry, such as missing photochemical mechanisms.

The most recently published HO<sub>2</sub> measurement uses a far-infrared spectrometer to observe thermal emission in a limb scanning mode from a balloon-borne gondola in the 49-23 km interval [Traub et al., 1990]. This measurement agrees with the column measurement of deZafra et al. [1984], leading to an HO<sub>2</sub> altitude profile from these two measurements in good agreement (within a factor of two) with theoretical models. In this paper we present new, high signal-to-noise, in situ HO<sub>2</sub> measurements in the 37-23 km interval that agree with the conclusions drawn by Traub et al. [1990] that HO<sub>2</sub> densities are not in profound disagreement with current models.

The instrument used here is based upon an OH detection technique. HO<sub>2</sub> is detected as above [Anderson et al., 1981] by adding NO to the atmospheric sample to convert HO<sub>2</sub> to OH. Nearly simultaneous measurements of the OH radical density are acquired during times when no NO is injected into the flow. The main body of balloon-borne OH measurements (see Stimpfle et al. [1989]) displays less scatter than the HO<sub>2</sub> data base and can be considered to be in broad agreement with model predictions. The new OH measurements reported here do not change that conclusion. However, by detecting OH and HO<sub>2</sub> simultaneously, the leverage provided by this technique to test HO<sub>x</sub> photochemical mechanisms used in model calculations is greatly enhanced. The measured ratio of HO<sub>2</sub> to OH provides a direct, in situ test of the small subset of reactions that control the rapid cycling of HO<sub>x</sub> radicals between the OH and HO<sub>2</sub> species. Additionally, the uncertainty in the ratio is less than that of the individual OH and HO<sub>2</sub> measurements, because error in the absolute OH detection sensitivity cancels out.

In this paper we discuss the experimental modifications made in preparation for the flight described here and present the results of the flight.

## Experimental

The OH detection instrument uses the LIF technique and has been described previously [Stimpfle et al., 1989; Stimpfle and Anderson, 1988]. OH is optically pumped in the A-X (1,0) band via the Q<sub>1</sub>(1) line at 282 nm with a 17-kHz repetition rate, Cu-vapor-laser pumped dye laser. OH A-X, (0,0) band fluorescence centered at 309 nm is observed with a filtered photomultiplier tube. Detection takes place in the core of a flowing sample of ambient

Copyright 1990 by the American Geophysical Union.

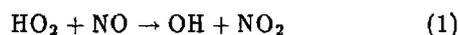
Paper number 90GL01825

0094-8276/90/90GL-01825\$03.00

stratospheric air as it is drawn through a cylindrical detection chamber during controlled balloon descent from peak altitude. The OH fluorescence signal,  $S^{\text{OH}}$ , is distinguished from the background signal by successively step tuning the laser on and off resonance with the OH absorption line in equal 10-s increments.

Two significant modifications have been made to the instrument in preparation for the flight reported here:

First, the ability to detect HO<sub>2</sub> in addition to OH was included by adding an NO gas addition subsystem to the detection pod. The



reaction converts ambient HO<sub>2</sub> to OH in the presence of excess NO. The NO is added to the sample flow through a circular array of 32, black teflon coated, 0.16 cm o.d., stainless steel tubes. The fixture is located 40.5 cm upstream of the fluorescence detection volume at a 12.7-cm i.d. section of the flow tube. This position typically allows 15 ms of reaction time at the flow velocities achieved in flight.

Second, the absolute OH detection sensitivity was increased by a factor of ~12 by adopting a new optical design for the fluorescence detection axis. In previous flights of this instrument (July 15, 1987 and July 6, 1988) the detector consisted of a fiber optic cable coupled to the curved entrance slit of a 1/4 m spectrometer. A PMT at the exit slit detected OH fluorescence. In order to increase the absolute OH detection sensitivity, the photomultiplier tube was relocated to the detection pod with a front-end optical system employing two optical elements. The first is a sealed quartz cell, 1-mm inner wall spacing, containing a solution of potassium hydrogen phthalate dissolved in methyl alcohol. This liquid, UV cut-off filter effectively absorbs Rayleigh ( $T \leq 1 \times 10^{-6}$  at 282 nm) and N<sub>2</sub> Raman ( $T \sim 5 \times 10^{-2}$  at 302 nm) scattered laser light while preserving a high transmission for OH fluorescence ( $T = 0.8$  at 309 nm). The second optical element is a 1.5-in diameter interference filter, with peak  $T = 0.22$  at 310 nm and a FWHM band pass of 4 nm. The optical elements are housed in a thermostated aluminum block to prevent temperature induced drifts of the interference filter pass band and transmission.

The instrument was calibrated before flight in the manner described previously [Stimpfle et al., 1989]. The observed count rate due to OH,  $S^{\text{OH}}$  (counts s<sup>-1</sup>), is given by the following expression:

$$S^{\text{OH}} = C^{\text{OH}} E Q [\text{OH}]$$

where  $C^{\text{OH}}$  is the calibration constant determined by direct measurement,  $E$  is the power (mW),  $Q$  is the fluorescence efficiency (unitless), and  $[\text{OH}]$  is the density (molec cm<sup>-3</sup>). The instrumental sensitivity was found to be:

$$C^{\text{OH}} = 1.4 \pm 0.5 \times 10^{-4} \left( \frac{\text{counts/s}}{\text{mW} \cdot Q \cdot [\text{OH}]} \right)$$

The flight NO injector head was tested in the flow bench and found to produce no observable losses of OH. The HO<sub>2</sub> conversion was tested qualitatively by checking that the selected NO mass flow rates were adequate for driving

reaction (1) to completion in the time scale allowed. The NO addition cycle is described below.

### Results and Discussion

The experiment was launched at 1602 UT (1023 CDT) on August 25, 1989 from the National Scientific Balloon Facility in Palestine, TX. A float altitude of 37.9 km was attained at approximately 1820 UT. The controlled descent period of the flight was initiated in the afternoon at 2202 UT, 51° solar zenith angle (SZA), and terminated at 2248 UT, 61° SZA, at an altitude of 23.1 km. The unexpectedly long period of time before descent was required by safety considerations concerning the location of the ground track defined by the balloon trajectory.

The instrument operated flawlessly during the 46 min descent period with excellent signal-to-noise. A 160-s segment of unaveraged (1/4 s), 309 nm detector data in the 31.5–30.7 km interval is shown in the top panel of Figure 1. The NO addition and laser tuning cycles are also displayed to illustrate their synchronization with the signal. The laser is tuned on and off resonance with the Q<sub>1</sub>(1) OH absorption line in the 20-s cycle shown in the lower panel of Figure 1. The NO addition cycle (middle panel) consists of two different, successively alternating, NO flow rates. The NO addition starts at the beginning of an off-resonant measurement to check for NO effects on the background. At the end of the following on-resonant measurement, the NO flow is terminated and immediately followed by a 4-s N<sub>2</sub> purge. The purge flushes NO out of the injector prior to the next on-resonant measurement, corresponding to an ambient OH measurement where there is no gas addition. NO addition is then repeated as above with a second NO flow, a factor of 0.4 that of the first. A complete 80-s cycle consists of four alternating OH and HO<sub>2</sub> measurements, with the HO<sub>2</sub> measurements made with two different NO flow rates. With a gondola descent velocity of ~5 m s<sup>-1</sup>, each radical measurement is completed within a vertical distance of 100 m.

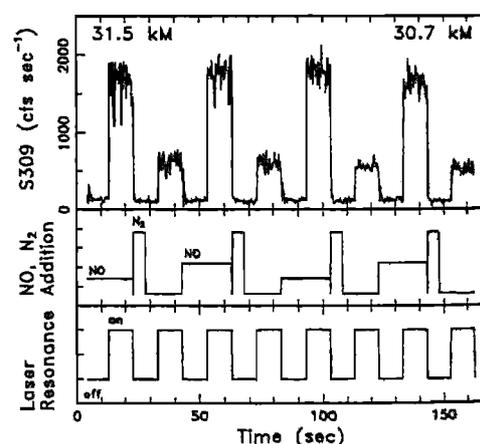


Fig. 1. Unaveraged 309 nm detector count rate data (top). NO and N<sub>2</sub> gas addition sequence (middle), and laser tuning cycle (lower) vs. time, during balloon descent through the 31.5–30.7 km interval.

The signal observed in the presence of NO is due to both ambient OH and that produced by NO reaction with HO<sub>2</sub>. Extracting HO<sub>2</sub> densities from the count rate data

is slightly complicated by the three-body reaction of OH with NO:



This reaction is slow at 37 km, but becomes increasingly important as air density increases. For this reason two different NO flow rates are used as a simple test to check completeness of the conversion reaction. The relative invariance of the observed signal to the two NO flow rates over a large dynamic range in pressure is evidence that the conversion was essentially optimal, i.e., in a region near peak conversion. Altitude dependent values for the fraction of ambient OH surviving recombination with NO ( $c_1$ ) and the fractional conversion of HO<sub>2</sub> to OH ( $c_2$ ) are calculated from the known temperature-dependent rate constants for (1) and (2) [DeMore et al., 1988], and the observed temperature and pressure.  $c_1$  decreases from 0.93 to 0.73 between 37 and 23 km, and  $c_2$  decreases from 0.92 to 0.70. Using these values, the HO<sub>2</sub> density is given by the following expression:

$$[\text{HO}_2] = ([\text{OH}]^{\text{NO on}} - ([\text{OH}]^{\text{NO off}} \times c_1)) \times 1/c_2$$

where  $[\text{OH}]^{\text{NO on}}$  is the observed OH in the presence of NO, and  $[\text{OH}]^{\text{NO off}}$  is the average, observed ambient [OH] bracketting  $[\text{OH}]^{\text{NO on}}$ . If we are not optimally converting HO<sub>2</sub> to OH we will systematically underestimate [HO<sub>2</sub>] by this method.

The observed OH and HO<sub>2</sub> profiles are shown in Figure 2. Each point is the result of a 20-s integration period, yielding a vertical resolution of ~100 m. Two HO<sub>2</sub> profiles are shown corresponding to the results obtained with each NO flow rate. There is an indication that the lower NO flow has slightly less conversion throughout the descent but the difference is not significant in comparison with the estimated uncertainty in the overall accuracy of  $\pm 35\%$  for [OH] given by uncertainties in the calibration and the fluorescence efficiency. The uncertainty in [HO<sub>2</sub>] is slightly greater,  $\pm 38\%$ , due mainly to uncertainty originating in the rate constant for (2). The increased OH sensitivity has greatly reduced imprecision due to counting statistics in comparison with previous flights. The shot-noise contribution ranges from  $\pm 0.7\%$  at 36 km to  $\pm 20\%$  at 23 km.

The observed fine structure in the OH and HO<sub>2</sub> profiles shown in Figure 2 cannot be considered real until we have eliminated possible instrumental effects as a cause. Photochemical mechanisms do not provide a clue concerning the authenticity of the observed structure in this case, because there is no significant covariance of OH and HO<sub>2</sub>. The observed scatter may simply represent our current instrumental precision, limited by our ability to account for frequency or power fluctuations, or small perturbations to the flow that are not indicated by the flow temperature measurement.

The HO<sub>2</sub> results, expressed in mixing ratio, are compared with previous HO<sub>2</sub> measurements in Figure 3. This figure is taken from the summary of Traub et al. [1990] in which the results of Helten et al. [1984] and Anderson et al. [1981] are scaled to matching SZA of 54° corresponding to the Traub FTIR measurement. Our results, measured between 51° and 61° SZA, need not be scaled

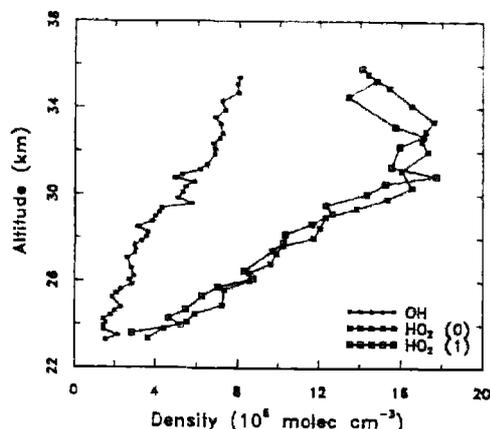


Fig. 2. [OH] and [HO<sub>2</sub>] vs. altitude observed Aug. 25, 1989, over a solar zenith angle variation of 51°–61°. HO<sub>2</sub> (0) and (1) refer to high and low NO flow rate results, respectively.

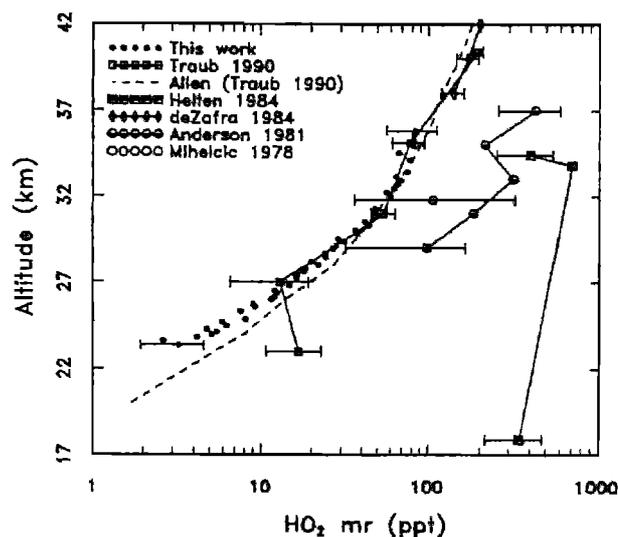


Fig. 3. Summary of HO<sub>2</sub> mixing ratio measurements vs. altitude from Traub et al. [1990], at 54° solar zenith angle. Measurements by Traub et al. [1990], Anderson et al. [1981], de Zafra et al. [1984], Helten et al. [1984], and model calculation by M. A. Allen (from Traub et al. [1990]). Amended to include results from Mihelcic et al. [1978] and this work.

for a meaningful comparison. As can be seen in Figure 3, this data set is in good agreement with the results of Traub et al. [1990] and therefore with the extrapolated profile of deZafra et al. [1984]. Convergence of HO<sub>2</sub> measurements obtained by three different techniques provides strong evidence that HO<sub>2</sub> has been validly measured in the stratosphere. Our HO<sub>2</sub> measurements will be repeated to further diagnose details of the NO conversion kinetics by modifying the gas addition scheme to include three NO flow rates in more rapid sequential addition.

The present OH results are compared with previous OH measurements made with this instrument in Figure 4. Since the earlier flights were characterized by relatively low SZA, 14°–9° [Stimpfle and Anderson, 1988] and 15°

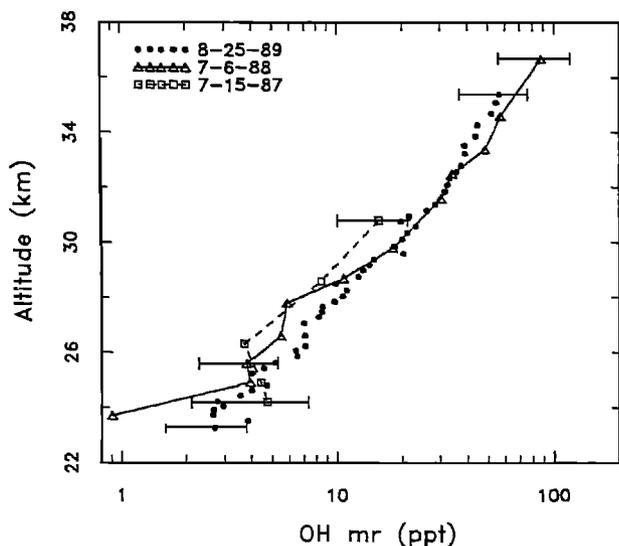


Fig. 4. Comparison of OH mixing ratio measurements vs. altitude made with this instrument, from Stimpfle and Anderson [1988], Stimpfle et al. [1989], and this work.

27° [Stimpfle et al., 1989], the present measurements have been scaled to ~ 24° to remove the SZA difference using typical, mid-latitude photochemical model results (S. Lloyd, private communication). The correction is a multiplicative factor ranging from 1.2 at 36 km to 2.0 at 23 km. The agreement among the three data sets is good if the simple standard of overlapping error bars ( $\pm 1\sigma$ ), which combine the precision and the estimated error in the absolute accuracy, is taken as the criterion. The most significant variation is the different shape of the July 6, 1988 and August 25, 1989 profiles in the 26–29 km region. Differences of this quality may simply represent year-to-year variation of dynamic and photochemical variables that control HO<sub>x</sub> source and sink species. Our simultaneous measurements of O<sub>3</sub> and H<sub>2</sub>O indicate that variation of these species is not responsible for the variation of the OH profiles in Figure 4. Variation of other parameters that can affect HO<sub>x</sub> species limits the level of detail to which comparisons of HO<sub>x</sub> data from different air masses can be carried out.

The major advance given by the data described here is that we are now able to examine two critical elements of HO<sub>x</sub> photochemistry with a data set from the same air mass. First, we can examine the ratio of HO<sub>2</sub> to OH that is believed to be controlled principally by NO, O<sub>3</sub> and O. This partitioning is critical in determining the interaction of HO<sub>x</sub> with the NO<sub>x</sub> and Cl<sub>x</sub> species in the stratosphere. Secondly, we can begin an examination of HO<sub>x</sub> production and loss balance. Production is proportional to [O<sub>3</sub>] and H<sub>2</sub>O mixing ratio throughout the stratosphere. Loss is proportional to the product of [OH] and [HO<sub>2</sub>] above 30 km and becomes dominated by reactions of OH with NO<sub>y</sub> species in the lower stratosphere. An examination of these

issues is presented in a separate paper [Wennberg et al., 1990].

**Acknowledgments** This flight could not have been carried out without the special assistance on short notice of a large number of individuals: At NASA headquarters, Bob Watson; at NSBF, Danny Ball and the NSBF support staff, which provided excellent assistance; and at Harvard, E. Weinstock, J. Demusz, N. Hazen, B. Heroux, T. Schiller, D. Spillane and E. Thompson. Funds provided by NASA contract NASW 3960.

#### References

- Anderson, J. G., H. J. Grassl, R. E. Shetter, and J. J. Margitan, HO<sub>2</sub> in the stratosphere: Three in-situ observations, *Geophys. Res. Lett.*, **8**, 289–292, 1981.
- de Zafra, R. L., A. Parrish, P. M. Solomon, and J. W. Barrett, A measurement of stratospheric HO<sub>2</sub> by ground-based millimeter-wave spectroscopy, *J. Geophys. Res.*, **89**, 1321–1326, 1984.
- DeMore, W. B., M. J. Molina, S. P. Sander, D. M. Golden, R. F. Hampson, M. J. Kurylo, C. J. Howard, and A. R. Ravishankara, NASA Public. 87-41, 1988.
- Helten, M., W. Patz, M. Trainer, H. Fark, E. Klein, and D. H. Ehhalt, Measurements of stratospheric HO<sub>2</sub> and NO<sub>2</sub> by matrix isolation and ESR spectroscopy, *J. Atmos. Chem.*, **2**, 191–202, 1984.
- Mihelcic, D., D. H. Ehhalt, G. F. Kulesa, J. Klomfass, M. Trainer, U. Schmidt, and H. Rohrs, Measurements of free radicals in the atmosphere by matrix isolation and electron paramagnetic resonance, *P. Appl. Geophys.*, **116**, 530–536, 1978.
- Stimpfle, R. M., L. B. Lapson, P. O. Wennberg, and J. G. Anderson, Balloon borne in-situ detection of OH in the stratosphere from 37 to 23 km, *Geophys. Res. Lett.*, **16**, 1433–1436, 1989.
- Stimpfle, R. M. and J. G. Anderson, In-situ detection of OH in the lower stratosphere with a balloon borne high repetition rate laser system, *Geophys. Res. Lett.*, **15**, 503–506, 1988.
- Traub, W. A., D. G. Johnson, and K. V. Chance, Stratospheric hydroperoxyl measurements, *Science*, **247**, 446–449, 1990.
- Wennberg, P. O., R. M. Stimpfle, E. M. Weinstock, A. E. Dessler, S. A. Lloyd, L. B. Lapson, J. J. Schwab, and J. G. Anderson, Simultaneous, in situ measurements of OH, HO<sub>2</sub>, O<sub>3</sub>, and H<sub>2</sub>O: A test of modeled stratospheric HO<sub>x</sub> chemistry, *Geophys. Res. Lett.*, this issue, 1990.
- WMO, *Atmospheric Ozone 1985*, Global ozone research and monitoring project, Report No. 16, Geneva, Switz., 1986.

J. Anderson, L. Lapson, R. Stimpfle and P. Wennberg, Harvard University, ESL, 40 Oxford St., Cambridge, MA 02138.

(Received July 12, 1990;  
accepted August 6, 1990)